

30 elements and their associated vulnerabilities; the consequences of the impacts; and the intensity of the
31 landslide event (Glade and Crozier, 2005). A landslide event may include more than one slope failure
32 triggered by the same phenomenon (e.g. a rainstorm). Interest in quantifying landslide risk has
33 developed since the attempt by the International Association of Engineering Geology (IAEG)
34 Commission on Landslides to compile a list of worldwide landslide events for the UNESCO annual
35 summary of information on natural disasters in 1971 (UNESCO, 1973). Although incomplete, five
36 years of records (1971-1975) recognised that landslides are a significant global hazard, with *c.*14% of
37 total casualties from natural hazards being attributed to slope failure (Varnes and IAEG Commission
38 on Landslides, 1984). Since, there has been a growing interest in landslide hazard and risk assessment
39 (Wu *et al.*, 2015).

40 Key elements of the assessment of landslide risk are coherent, high quality landslide
41 databases and inventories (van Western *et al.*, 2008; Van Den Eeckhaut and Hervás, 2012; Taylor *et*
42 *al.*, 2015). Inventories provide systematically compiled lists of landslide events that have occurred
43 over a specific spatial scale (*e.g.* within a nation) within a set period of time, or that result from a
44 single, catastrophic triggering event (Hervás and Bobrowsky, 2009). Databases organise inventory
45 information so that it is structured and searchable. Spatio-temporal analysis of global records of
46 landslides have demonstrated the extent to which landslides impact on society, and have identified
47 geographical regions and countries most exposed (Petley, 2012). Several different global databases
48 are actively maintained (*e.g.* the EM-DAT International Disaster Database, The NASA Global
49 Landslide Catalogue, and the Global Fatal Landslide Database (GLFD) on which this study is based),
50 and their merits and limitations are discussed by Van Den Eeckhaut and Hervás (2012) and
51 Kirschbaum *et al.* (2015). Global disaster databases are also maintained by risk reinsurers, however
52 landslides are often included within broader categories (such as geophysical hazards or within
53 weather-related hazards), and the majority of data are not freely available.

54 Relative to other natural disasters, the International Disaster Database (EM-DAT) suggests
55 that landslides account for 4.9 % of all natural disaster events and 1.3% of all natural hazard fatalities
56 between 1990 and 2015; 54% of these landslide events occurred in Asia (Guha-Sapir *et al.*, 2018).

57 However, the dedicated global landslide databases indicate that global multi-peril databases
58 underestimate the impact of landslides on society. Petley (2012) showed that the EM-DAT database
59 underestimated the number of fatal landslide events by ~2000% and fatalities by 430% between 2004
60 and 2010, whilst Kirschbaum *et al.* (2015) showed that the EM-DAT database underestimated the
61 number of fatal landslide events by ~1400% and fatalities by 331% between 2007 and 2013. For the
62 most-part this under-reporting is associated with the perception of landslides as a secondary hazard,
63 with the cause of death often being recorded in connection with the primary hazard (e.g. an
64 earthquake rather than a co-seismic landslide), rather than the actual cause of the loss.

65 Past studies on global landslide distribution have focused on rainfall-triggered events,
66 recognising the importance of rainfall and climate in inhabited regions with steep slopes (Dowling
67 and Santi, 2014; Kirschbaum *et al.*, 2012; 2015). This paper not only provides a key update on the
68 impact of landslides worldwide, extending Petley (2012) to include landslides from 2004 to 2016, the
69 study considers trends in landslides triggered by human activity. Thereby, adding to the discussion on
70 climate versus human disturbance as current and future drivers of landslide incidence (Crozier, 2010).

71 **2. The Global Fatal Landslide Database**

72 The GFLD (formerly termed the Durham Fatal Landslide Database) has been compiled using
73 systematic, English language based, metadata search tools that identify relevant reports of landslide
74 activity (including all mass movements falling within the definition of Hungr *et al.* (2014) on a daily
75 basis (Petley *et al.*, 2005; Petley, 2010; 2012). In common with other hazard databases (Tschoegl *et al.*
76 *et al.*, 2006; Taylor *et al.*, 2015), mass media reports provide a first alert for fatal landslide occurrence
77 and impact. Reports are corroborated and data updated by source triangulation using government and
78 aid agency reports, academic papers and personal communications, as new information becomes
79 available. The dataset has been consistently collected and managed since 2004, following a period of
80 methodological development between 1 September 2002 and 31 December 2003 (Petley, 2012). The
81 approach is differentiated from that of Kirschbaum *et al.*, (2010; 2012; 2015) because: (1) only
82 landslides that cause loss of life are included; and (2) all landslides are included, as opposed to only
83 those triggered by rainfall. In addition, the GFLD has been compiled over a longer period. Although

84 media reporting tends to be biased towards landslides with human casualties (Carrara *et al.*, 2003),
85 which is favourable for a database of this nature, it is recognised that the data collected is to some
86 degree an underestimate of the number of fatal landslides, and their associated losses. Landslides that
87 occur in remote mountain regions, or that result in a small number of fatalities, are less likely to be
88 reported than multi-fatality landslides and/or those that occur in urban centres (Petley, 2009).
89 Reliability of reporting is also spatially variable, based on the robustness of regional communication
90 networks, which are considered more consistent in developed nations (Petley, 2010; Kirschbaum *et*
91 *al.*, 2010), and in some cases political considerations (*e.g.* very few landslides are recorded in North
92 Korea). The true number of fatalities may be slightly underestimated when victims die of landslide
93 derived injuries weeks to months following the event (Petley, 2012). Furthermore, solely non-English
94 reporting of landslides will account for some missed reporting. However, Sepúlveda and Petley
95 (2015) compared the GFLD with an independently compiled database based on original Spanish and
96 Portuguese language reports for Latin America, and found a difference of only 5% of total records,
97 generally associated with landslides with small numbers of fatalities. Combined these effects may
98 underestimate the true level of loss by up to 15% (Petley, 2012); however the methodology of
99 collation of the GFLD is considered robust.

100 Since 2004, the database has been compiled to include the date of occurrence, the description
101 of landslide location; an approximate latitude and longitude for that location; the country and
102 geographical region (based on UN classifications, UNSD, 2018) in which the landslide occurred; the
103 number of fatalities and injuries; and whether the event was triggered by rainfall, seismicity or
104 another cause. Seismically triggered landslides in the database are excluded from analysis herein,
105 because the catalogue of events is not considered complete (see Petley, 2012). These equate to 168
106 earthquake events and 3978 fatalities. In preparation of this paper, all landslide reports were reviewed
107 to enhance the classification of the trigger event according to Table 1, using keyword searches in the
108 original text describing the landslide. The description of the landslide event location may be specific
109 to a section of road or village, or give a more general location within an administrative division (such
110 as a county, or state). The location of all landslide events are known within political country

111 boundaries. To estimate the spatial precision of each landslide report, location descriptions were
112 related to spatial databases of administrative boundaries (GADM, 2017), transport network maps
113 (Google Maps, 2018; Open Street Map, 2018), and in some cases individual landslides could be
114 identified from satellite imagery (Google Earth, 2018; Planet Team, 2017). For administrative units
115 such as villages or states, polygon area from GADM (2017) provided the measurement of spatial
116 precision. For stretch of road, a polyline of the road length was created using transport network data
117 (Open Street Map, 2018) and a 500 m buffer applied; the area of the buffer provided the precision
118 estimate. The median spatial precision of entries is 681 km², with an interquartile range of 1 to 3477
119 km². A sample of the database and further information are provided in Supplementary Material.

120

121 3. Global Fatal Landslide occurrence, 2004 to 2016

122 The total number of fatal landslide events recorded worldwide, excluding those triggered by
123 earthquakes, over the twelve calendar years between 2004 and 2016 (inclusive) was 4862. The
124 majority of events (95%) involved a single slope failure. The spatial distribution of landslides (Fig. 1a
125 and 1c) is clearly heterogeneous, with high areas of incidence in:

- 126 • Central America between Costa Rica and the South of Mexico
- 127 • The Caribbean islands.
- 128 • South America, along the Andes mountain range from Venezuela to Bolivia and to a
129 lesser extent Chile, with another cluster of events on the east coast of Brazil around
130 the states of Sao Paolo and Rio de Janeiro.
- 131 • East Africa, around the borders between Tanzania, Rwanda, Burundi, Kenya, Uganda
132 and Democratic Republic of the Congo.
- 133 • Asia, which is the site of the highest number of events (75% of landslides).
134 Substantial numbers of landslides occur along the Himalayan Arc, in states across
135 India and southeast China, as well as high numbers in the neighbouring countries of

136 Laos, Bangladesh and Myanmar, and southwards on islands that form Indonesia and
137 the Philippines.

- 138 • There are smaller clusters in Turkey and Iran, as well as in the European Alps.

139

140 Fatal landslide events cluster around cities (Fig. 1c) and occur most frequently in countries
141 with lower Gross National Income (GNI, Fig. 1c) at locations known to be susceptible to landslides,
142 based on the analysis of physical characteristics of the environment (see Hong *et al.*, 2007; Stanley
143 and Kirschbaum, 2017). Textual analysis of landslide reports shows many events occurred in mines or
144 quarries (423 landslides), and 568 landslides in the dataset occurred on roads. Relative poverty is also
145 emphasised in reporting: the term ‘slum’ is explicitly used to describe the impacted community 29
146 times, while broader terms to indicate relative poverty are used 267 times within landslide reports.
147 These observations support previous research that fatal landslides are most prevalent in densely
148 occupied urban centres (Alexander, 1989; Anderson, 1992; Petley, 2009), along roads (Hearn, 2011;
149 Lee *et al.*, 2018), and at sites rich in natural resources (Zou *et al.*, 2018). In common with other
150 natural hazards, the poor are disproportionately affected by landslides (Hallegatte *et al.*, 2016).

151 Fig. 2 shows landslide occurrence in pentads, smoothed with a 25-day (*i.e.* five pentad)
152 moving average. The most landslide events in a single pentad was 48, in early October 2009; of these
153 45 were triggered in a single day (8th October 2009) by Typhoon Parma in the Philippines. Rainfall is
154 the leading trigger of landslides. The majority of non-seismic fatal landslide events (2004-2016) in the
155 database were triggered by rainfall (79%). Fig. 3a shows landslide events triggered by rainfall in
156 pentads, compared with the complete non-seismic landslide event dataset. The data series are strongly
157 correlated ($R, 0.933, p\text{-value}, 0$), indicating that rainfall triggered landslides explain 93% of the
158 variance of the complete dataset. Fig. 3b shows landslide events that were not triggered by rainfall,
159 and where the trigger is known (*e.g.* mining). We term these events “non-seismic non-rainfall
160 triggered (NSNR) landslides” herein. These landslide events constitute 16% of the complete dataset
161 and present a different pattern through time when compared with rainfall-triggered landslide events.

162 There is a notable increase in the number of landslide events with NSNR triggers from about 2006,
163 which we ascribe to improved event capture.

164 The rainfall-triggered landslide data in Fig. 3a (and the complete landslide series, Fig. 2)
165 contain a strong seasonal pattern of landslide occurrence through the annual cycle, as noted by Petley
166 (2012). Autocorrelation measures the linear relationship between lagged values of a time series. The
167 autocorrelation of the rainfall-triggered pentad landslides series (Fig. S1) shows the correlation
168 coefficient between the original series and a lagged version of the series, where the series lags
169 between 1 to 948 pentads (5 days to ~13 years). The autocorrelation oscillates around 73.5 lags
170 (pentads), equating to one calendar year. This pattern is indicative of annual seasonality in the data.
171 Conversely, the autocorrelation of the NSNR landslides pentad series (Fig. S2) does not contain this
172 pattern and the correlation coefficients are generally weak. This indicates that there is no seasonal
173 pattern in the NSNR landslide series, which is to be expected in events that are not triggered by
174 meteorological processes.

175

176 **3.1. Seasonality**

177 Landslide event occurrence peaks in the northern hemisphere summer, and there is notable inter-
178 annual variation, both in the size and shape of the annual cycle. Seasonality in the global series (Fig. 2
179 and 3a) is associated with the annual cycle of rainfall-triggered landslides in South, South East and
180 East Asia, and South and Central America (Fig. 4). Combined, these geographical regions contain
181 88% of all rainfall-triggered landslide events and account for 96% of variance in the global seasonal
182 cycle (Table B1). There is a correlation between the mean monthly rainfall (data from GPCC, 2018;
183 Xie *et al.*, 2013) and landslide series, for four of five regions (Fig. 5 and Table 2), reflecting the
184 triggering effect of seasonal rainfall. However, the strength of relationship between seasonal patterns
185 of rainfall and the seasonal pattern of landslide events is variable between regions. The pattern is
186 strongest in East Asia and South Asia. This corroborates Petley (2012) who identified the strong

187 relationship between landslide occurrence and seasonal rainfall from a shorter period of data (2004 to
188 2009).

189 Seasonal rainfall in East and South Asia is associated with the onset and withdrawal of the
190 Asian monsoon (*e.g.* Webster, *et al.* 1998), delivered by the seasonal reversing of winds to flow from
191 ocean to land in the summer months, resulting in the majority of annual rainfall occurring between
192 June and September (Turner and Annamalai, 2012). In South Asia, landslide incidence increases in
193 Nepal, India, Bangladesh, Bhutan and northern Pakistan during the summer monsoon. India and
194 Nepal contribute 16% and 10% respectively of all rainfall-triggered landslide events in the global
195 dataset; of these 77% and 93% occurred during the summer monsoon, meaning 21% of all rainfall-
196 triggered landslide events globally were triggered by seasonal monsoon rainfall in India and Nepal. In
197 East Asia, tropical cyclones extend the length of the rainfall season: 109 landslide events were
198 triggered by typhoons between April and October in China, Japan and South Korea, representing 16%
199 of rainfall-triggered landslide events in East Asia, and 3% of global rainfall-triggered landslide events.
200 The East Asia landslide record is dominated by events in China (81%, 503 landslides), of which 409
201 landslide events were triggered during the summer monsoon rainfall season. China alone contributes
202 15% of all global rainfall-triggered landslide events, although the pattern is heterogeneous.

203 Although, the seasonal landslide series for Central and South America do not explain much
204 variance in the global seasonal landslide cycle (because of the comparatively low number of
205 landslides), there is strong correlation between patterns of landslides in the region and patterns of
206 rainfall (Table 2). Central America and parts of the Caribbean experience a summer rainy season
207 between May and October, associated with the position of the Inter-Tropical Convergence Zone
208 (ITCZ; Garcia *et al.*, 2009). The season is bimodal, with peaks in rainfall on either side of a
209 midsummer drought, between late June to August (Magaña *et al.*, 1999). The season is enhanced by
210 the Atlantic basin hurricane season from 1 June to 30 November (NOAA, 2018a). The pattern of
211 landslides reflects these rainfall drivers. South America spans ~70° of latitude leading to local
212 variability in climate (Sepúlveda and Petley, 2015). The peak annual rainfall for the continent as a
213 whole occurs during the period from December through to February, delivered by the South American

214 Monsoon System (SAMS), which is driven by the position of the ITCZ to the south of the equator
215 (Garcia *et al.*, 2009). However, in parts of south-eastern Brazil, where there is a prevalence for fatal
216 landslides (Fig. 1), the rainy season extends into March (Rao and Hada, 1990). In northern Peru,
217 rainfall peaks between April and June in the west, and is bimodal in the east, with peaks in April and
218 December (Espinoza *et al.*, 2009). Colombia's meteorology is particularly complex due to the
219 convergence of the Equatorial Mid-tropospheric Easterly Jet (EMEJ) and the Choco Jet; the resulting
220 rainfall distribution is bimodal, with peaks in April-June and August-September, depending on precise
221 location and the choice of rainfall data and model (Sierra *et al.*, 2015). Most rainfall-triggered fatal
222 landslide events in South America occur in Brazil (37%) and Colombia (32%), most notably in south-
223 eastern Brazil and central Colombia, and this is evident in the distribution of annual rainfall and
224 landslide occurrence (Fig. 5d).

225 The weak relationship between rainfall and landslides in South East Asia reflects the complex
226 weather systems operating in the region. Most landslide events occurred in the Philippines (46%) and
227 Indonesia (32%). Typhoons caused 22% of rainfall-triggered landslide events in the region, and 5%
228 globally; most typhoon-triggered landslide events occurred in July through to October (75%), in line
229 with the main tropical cyclone season. In the Philippines, 42% of rainfall-triggered landslide events
230 were caused by Typhoons, whilst the equivalent value for Vietnam was 22%, although of a much
231 lower total. The pattern of monsoon rainfall in Indonesia and the Philippines varies by geographical
232 location. In the west of the Philippines, summer monsoon occurs between June and October, while in
233 the east, the winter monsoon occurs between October and March (Kubota *et al.*, 2017). This pattern is
234 evident in the distribution of rainfall-triggered landslides in the Philippines (Fig. 1a). The onset and
235 termination of the monsoon in Indonesia varies from September to June in north Sumatra and late
236 November to late May in east Java (Naylor *et al.*, 2007). Consequently, 72% of rainfall-triggered
237 landslide events occur between November and April, when the majority of Indonesia is experiencing
238 monsoon rainfall. The peak in landslide activity relative to rainfall in August to October in South East
239 Asia (Fig. 5b) is mainly due to the localised typhoon rainfall not captured in the regional rainfall
240 average.

241 **3.2. Medium term trend in landslide occurrence**

242 There was a general increase in recorded landslide occurrence between 2004 to March 2010, followed
243 by a general decrease in landslide occurrence through to April 2015, after which landslide incidence
244 has generally increased (Fig. 6a). Petley (2012) identified improvements in the reporting of single
245 fatality landslides as contributing to the general increase in events in the fatal landslide record from
246 2004 to 2010. The number of fatalities resulting from non-seismic landslide events between 2004 and
247 2016 was 55997. Fig. 6b, shows that the pentad series of fatality is very noisy; the data do not contain
248 an increasing or decreasing trend, nor are there distinguishable medium-term peaks in the data. Very
249 few landslide events generated more than 1000 fatalities (0.1%), and only one landslide resulted in
250 more than 5000 fatalities. This was the Kedarnath landslide in June 2013 in Uttarakhand state, India,
251 which was caused by extreme meteorological conditions that generated flooding and two large
252 landslides in a mountainous area occupied by thousands of religious pilgrims (Allen *et al.*, 2016).

253 Landslide events by the number of fatalities are grouped by the infinite series (1, 2, 4, 8,
254 16...). There is a significant increasing trend in single-fatality landslides (Fig. 6c); 29% of landslides
255 were single-fatality events. There is also a weaker decreasing trend in landslide events resulting in 64
256 to 128 fatalities (Fig. 6d); 1% of landslide events were in this group. No other grouping contained a
257 significant trend with time. Both the single fatality and 64 to 128 fatality series are above the
258 regression line in 2010 (Fig. 6c and 6d). Removing these two groups from the global series (Fig. 6e) it
259 is evident that single fatality events enhanced the peak around 2010, and in 2016.

260 By year, different geographical regions experience above/below average landslide activity
261 (multi-fatality landslide events, Fig. 6f; single-fatality landslides, Fig. 6g). In 2005, 2009, 2010 and
262 2011, several regions experienced greater than average landslide occurrence simultaneously. The high
263 impact of landslides globally in 2010, has been discussed by previous authors (Kirschbaum *et al.*,
264 2012; 2015; Petley, 2012; Sepúlveda and Petley, 2015). The peak in landslide activity was generated
265 by anomalous landslide occurrence in several regions simultaneously (Fig. 6f and 6g), but overall the
266 geographical pattern of rainfall triggered landslide events in 2009 and 2010 reflects the occurrence of
267 a moderate El Niño in 2009 and a moderate La Niña in 2010 (NOAA, 2018b).

268 In Central America, Kirschbaum *et al.* (2012) showed that rainfall was significantly above
269 average in the summer months in 2010, particularly in September. This increase was linked to the
270 known impacts of La Niña events on tropical cyclone frequency and track (*e.g.* Elsner *et al.*, 1999;
271 Curtis *et al.*, 2007). By number, 2010 was the year in which the most landslides (17 events, compared
272 with an average 6 events per year), were directly associated with tropical cyclones in reports or related
273 to storm tracks (based on NOAA, 2018c). Although, these landslide events only equate to 35% of all
274 rainfall-triggered landslide events within 2010, the remaining 65% of events, not triggered by a
275 tropical cyclone all occurred during the hurricane season (May to November), likely due to unsettled
276 weather associated with warm sea surface temperatures in the region. Central America receives
277 tropical cyclones from the Atlantic basin and the North Pacific basin (NOAA, 2018c). Storms from
278 the Atlantic basin may make landfall along the eastern coastline of Central America and travel inland,
279 occasionally retaining enough energy to cross over into the Pacific. Storms that have crossed over
280 basins or new storms, which have formed in the Northeast Pacific basin, may make landfall on the
281 western coast of Central America. Not only were the frequency of landfalling tropical storms and
282 hurricanes elevated from both basins in 2010, but the track of these storms intercepted populated areas
283 in steep terrain (NOAA, 2018c). The majority of rainfall-triggered landslide events in Central
284 America in 2010 were in Mexico and Guatemala (43% and 37%, respectively). In Guatemala, eight
285 landslide events were triggered by Tropical Storm Agatha in late May 2010, causing 182 fatalities.
286 Four landslide events were associated with Hurricane Alex which travelled up the east coast of
287 Guatemala, Honduras and then inland to Mexico in late June-July 2010. Hurricane Karl then made
288 landfall on the east coast of Mexico in September: two landslide events are associated with this storm
289 (killing 12), but a succession of fatal landslides in the states of Oaxaca, Chiapis and Puebla through
290 which the hurricane passed, were noted in the weeks following the storm.

291 Sepúlveda and Petley (2015) observed a weak correlation between La Niña conditions in late
292 2010-2011 and heightened landslide activity in Colombia and Venezuela. Considering a longer time
293 series (2004 to 2016), this study identifies above average landslide activity in several nations in South
294 America in 2009 and 2011. In Brazil, 54% of all rainfall-triggered events occurred between 2009 and

295 2011. Activity peaked in December 2009 to April 2010 (El Niño) and January 2011 (La Niña),
296 corresponding with the seasonal El Niño-Southern Oscillation (ENSO) rainfall patterns observed by
297 Grimm and Tedeschi (2009). The number of landslide events in Venezuela and Colombia between
298 2009 and 2011 peaked in November 2010, associated with positive rainfall anomalies during the
299 austral summer La Niña (Tedeschi *et al.*, 2013).

300 The majority of landslide events in East Asia occur in China (83%); in 2010, 87% of all
301 rainfall-triggered events were located in China, and rainfall-triggered landslide occurrence (67
302 landslide events) was above the mean (45 landslide events). From a shorter period of observation,
303 Kirschbaum *et al.* (2012) identified a high incidence of rainfall-triggered landslides (fatal and non-
304 fatal) in central eastern China in 2010: particularly in July and August corresponding with a peak in
305 rainfall. Rainfall-triggered landslides were above average for most months in 2010 in China, but May
306 through to September was very active (57 landslide events compared with an average 38). The East
307 Asian subtropical summer monsoon (a component of the East Asian monsoon) has a significant effect
308 on seasonal variations in rainfall across China (He and Liu, 2016), and rainfall patterns alter in
309 response to ENSO conditions (Yang and Lau, 2004; He *et al.*, 2007; Zhou *et al.*, 2014).

310 In China in 2010, there were fewer than average landslide events triggered by tropical
311 cyclones from the Northwest Pacific basin. There was low typhoon activity due to the rapid transition
312 from the 2009/2010 El Niño to the 2010/2011 La Niña, which altered airflows in the North West
313 Pacific basin (Kim *et al.*, 2012). Conversely, in the Philippine domain, tropical cyclone occurrence
314 was above average in July to December 2009 (Corporal-Lodangco *et al.*, 2015). During the northern
315 hemisphere summer months of an El Niño, the genesis location of tropical cyclones shifts eastwards
316 (Chan 1985; 2000; Chia and Ropelewski, 2002). In these conditions, cyclones travel further before
317 they may make landfall, enabling them to strengthen (Camargo and Sobel, 2005), and there is a
318 tendency for more storms to affect the northern-central Philippines (Lyon and Camargo, 2009). In
319 2009, 67% of rainfall-triggered landslide events in the Philippines were associated with tropical
320 cyclones: 60 landslide events compared with an average 12 triggered by tropical cyclones. As noted
321 previously, many of these were triggered on the same day (8 October 2009) by Typhoon Parma.

322 Although the peak in landslides in South East Asia in 2009 is dominated by typhoon triggered
323 landslides in the Philippines, there was an increase in landslides in Indonesia (33 landslide events
324 compared with an average of 24 per year); of these 24 events were triggered by rainfall, 8 by mining
325 and one trigger was not known. Rainfall-triggered landslide events were very slightly above average
326 in Indonesia in 2009 but it was the events triggered by human activity that contributed most to the
327 anomalous landsliding in Indonesia. These landslides are discussed in the next section.

328 Between 2004 and 2016, four El Niño events occurred: weak El Niño (2004/2005,
329 2006/2007), strong El Niño (2009/2010) and very strong El Niño (2014/2016; NOAA, 2018b). Weak
330 La Niña was observed in 2005/2006, 2008/2009, 2016, and strong La Niña in 2007/2008 and
331 2010/2011 (NOAA, 2018b). There does not appear to be a consistent relationship between ENSO
332 phase and the regional distribution of landslides, although elevated regional rainfall (and thus
333 landslides) has been associated with ENSO sea-surface temperature (SSTs) anomalies. The peak in
334 landslide events in Central America in 2005 is composed predominantly of tropical storm and
335 hurricane triggered landslides in El Salvador, Mexico, Guatemala and Honduras. The 2005 North
336 Atlantic Hurricane season was the most active since records began in 1851, driven by high SSTs in
337 the Tropical North Atlantic (10°-20°N) linked with global warming and the 2004/2005 El Niño
338 (Trenberth and Shea, 2005). Landslide events were also above average in 2005 in East Asia: most
339 events occurring in China, triggered by monsoon rainfall. In South Asia, landslide events peaked in
340 2007, 2014 and 2016, the majority associated with monsoon rainfall in Bangladesh, India, Nepal and
341 Pakistan. Variability in rainfall from the South Asian monsoon is related to the interaction between
342 SSTs in the Indian Ocean dipole (IOD) and ENSO (*e.g.* Ashok *et al.*, 2007; Lu *et al.*, 2017).

343 The complexity of climate systems means it is not possible to draw conclusions on the
344 relationship between climate mode and landslide occurrence from this 13 year global dataset.
345 However, longer local records show promise at unpicking the impact of climate cycles on landslides.

346 **3.3.NSNR landslide Triggers**

347 Of the 4862 non-seismic landslide events in the complete database, 770 (16%) were generated by a
348 NSNR trigger, and resulted in a total 3725 fatalities (Fig. 7). The majority of landslides were triggered
349 by mining (232 multi-fatality landslide events, 67 single-fatality landslides), construction (170 multi-
350 fatality landslide events, 140 single-fatality landslides) or illegal hillcutting (60 multi-fatality
351 landslide events, 27 single-fatality landslides); and the majority of fatalities in all cases were people at
352 work (90%, 76% and 84% respectively). Globally there is a statistically significant increase in events
353 by these three triggers (Fig. 8a, 8b and 8c); multi-fatality landslide events are differentiated from
354 single-fatality landslides, which increased with time independent of trigger (Fig. 6c). By country,
355 most construction triggered landslide events occurred in India (28%), followed by China (9%),
356 Pakistan (6%), the Philippines (5%), Nepal (5%) and Malaysia (5%; Fig. 9a). On average construction
357 triggered landslide events killed three people, however a particularly severe landslide in Shenzhen,
358 China in December 2015 killed 77 people. The event involved the collapse of construction waste on
359 worker quarters in an industrial site. Interestingly, the context in which the landslides occur differs
360 between countries. In China, the majority of events (52%) occur in urban construction sites, while
361 very few landslides occur on roads (7%). Conversely, in India and Nepal, 30% and 43% of landslide
362 events triggered by construction occurred on roads.

363 Transportation is a “crucial driver of development” (World Bank, 2018b); however, in
364 mountain regions roads are closely connected with landslide risk (Lennartz, 2013). The road network
365 in Nepal has quadrupled in length over the last 18 years (Govt. of Nepal, 2016), and in India it has
366 nearly tripled in length in 24 years (Govt. of India, 2016). Population growth is frequently
367 accompanied by the expansion of infrastructure and settlements (Gardner and Dekens, 2007), and this
368 is true in India and Nepal, which have grown by ~7% between 2010 and 2015 (World Bank, 2018a).
369 Both countries are on a trajectory to expand their national road networks further. Increased landslide
370 activity in the Himalayan region has been associated with road construction (Ives and Messerli, 1989;
371 Haigh *et al.*, 1989; Valdiya, 1998; Barnard *et al.*, 2001; Petley *et al.*, 2007; Sait *et al.*, 2011; Singh *et*
372 *al.*, 2014). Hearn and Shakya (2017) highlighted that road construction without proper route choice,
373 engineering design and management of spoil, increases landslide susceptibility. Fatal landslides

374 triggered by road construction indicate that excavation may not always be undertaken with due care
375 and appropriate slope engineering. Furthermore, the coincidence of construction worker and road user
376 fatalities from the same landslide suggests that there is pressure to keep roads under construction
377 open. Ives and Messerli (1989) emphasised the economic impact when roads are closed.

378 Between 2004 and 2016, China experienced a 6% growth in population to 1.379 billion, and a
379 16% rise in the proportion of the population living in urban areas (World Bank, 2018a). Urban growth
380 in China is driven by political policy for economic growth; economic reforms from 1978, opened
381 China's markets to foreign investors and relaxed migration controls, prompting rapid rural-urban
382 migration (Ma, 2002; Anderson and Ge, 2004). Although urbanization is encouraged by China to
383 increase domestic consumption, urban growth is often uncontrolled (Fang and Pal, 2016), leading to
384 rapid land conversion, dispersion and fragmentation of development (Schneider and Woodcock,
385 2008). Critically, many of China's largest cities are bounded by mountains, and urban sprawl is
386 encroaching on land unsuitable for development (Yu and Li, 2011). Reports in the database indicate
387 that fatal landslides in urban construction sites in China often occurred when engineered cut slopes
388 failed above the construction site (*e.g.* Zhang *et al.*, 2012), from improper construction of foundations
389 leading to building collapse before completion (*e.g.* Srivastava *et al.*, 2012), or from mismanagement
390 of construction and demolition waste (*e.g.* Yang *et al.*, 2017). In these entirely preventable
391 circumstances, explicit national regulation and enforcement should reduce construction related
392 landslide impact in China.

393 The increase in events triggered by mining is driven by the increase in landslides triggered by
394 illegal or unregulated extraction (Fig. 8d); landslides triggered by legal mining (Fig. 8e) or where the
395 legitimacy of the mining is unknown (Fig. 8f), do not show a statistically significant trend. By
396 country, India (12%), Indonesia (11.7%), China (10%), Pakistan (7%) and Philippines (7%) contribute
397 most to the record of landslides triggered by mining (Fig. 9b). Fatal landslides triggered by illegal
398 mining practises have occurred in 32 countries (Fig. 9c). By number of events, Indonesia (24) and
399 India (15) rank the highest, however by number of fatalities Myanmar (403 fatalities from 9 landslide
400 events) stands out. Shifts in spending power and the infusion of the internet and smart-technology in

401 daily life have driven an exponential increase in the consumption of electronics, placing pressure on
402 the demand for rare earth elements (Dutta *et al.*, 2016). Furthermore, growth in the precious stone
403 market fuelled both by economic uncertainty, and a growing middle class in Asian nations such as
404 China, where gemstones are a key part of cultural heritage (The Economist, 2011), is thought to have
405 led to an increase in the number of small-scale mining operations globally (Hruschka and Echavarría,
406 2011), and the upscaling of small-scale mines to larger scale operations. Fatal landslides in Myanmar
407 (Burma) have significantly increased because of the unregulated expansion in jade mining within the
408 Kachin state. Critically, the high value of jade and lack of enforced operator accountability appears to
409 be driving poor mining practises, which place workers and local residents at risk of slope collapse
410 (Global Witness, 2015). Demand for rare earth elements and gemstones are thus driving an increase in
411 mining-related landslides, with the potential for landslide occurrence to rival that associated with rural
412 road expansion.

413 Cutting slopes for the purposes of obtaining earth surface materials, or to alter slope geometry
414 during construction, may result in slope failure if the site is not properly engineered. The term hill-
415 cutting is used here in relation to discrete slopes that have been altered without permission for the
416 purposes of small-scale construction, earth material extraction or agriculture. Hill-cutting is most
417 strongly associated with urban areas in Bangladesh in the academic literature (*e.g.* Chittagong;
418 Ahmed, 2015 or Syhlet; Islam *et al.*, 2006). In the fatal landslide database it is an increasing problem
419 in Bangladesh, India and Nepal (Fig. 8c and 9d). Most fatalities occurred as people collected hillslope
420 materials for construction of their housing in rural communities, and reports indicate those involved
421 were from poor families living in informal settlements. In total, 11 of the 87 landslide events were
422 directly related to the practice of using hillslope coloured clay for the decorative coating of houses for
423 a religious festival; of these, nine occurred in Nepal. Critically, children are often caught up in slides
424 triggered by hill-cutting in Nepal: at least 40% of landslide victims were children, while a further 25%
425 of victims were a combination of adults (predominantly women) and children working together.
426 Conversely, in Bangladesh the majority of victims were adults (78%) of which 79% were male. In
427 Nepal, India and Bangladesh, clay is an important local building material for housing, particularly in

428 settlements not connected to the road network, but there is a legal framework in Bangladesh to
429 prevent hill-cutting (Building Construction Act 1952, 1990 and the Bangladesh Environmental
430 Conservation Act 1995; Murshed, 2013). Building codes in Nepal provide basic guidance on slope
431 stability, specifically slope excavation, identification of slope instability and construction of
432 foundations (DUDBC, 1994); however residents in rural communities may not have access to this
433 information and be unaware of the hazard (Oven *et al.*, 2008). Furthermore, in India it was noted that
434 building regulations do not account for the geo-environmental context of the settlement, sometimes
435 lack clarity, and are difficult to uphold due to a shortage of technical experts and inadequate provision
436 to stop illegal activity (Kumar and Pushplata, 2014).

437 While this section discusses fatal landslides triggered by human activity, many rainfall-
438 triggered landslides occur on slopes which have been modified during construction (82 landslide
439 events), agriculture and forestry (45 landslide events) and mining (123 landslide events); or at sites
440 where storage of waste has not been poorly managed (16 landslide events). Of course, it is expected
441 the majority of fatal landslides (94%) will occur within settlement boundaries or along infrastructure,
442 however it is evident from this database of events that human action damages slopes increasing their
443 susceptibility to fail.

444 **4. Discussion and Conclusion**

445 With the benefit of a 13 year time series, this study builds on past analyses of the GLFD, providing
446 not only an update on the spatial and temporal distributions of landslide impact, but also serving to
447 highlight the importance of annual climate variability in specific landslide prone regions on the global
448 record. In addition, it provides new insights into the impact of human activity on landslide incidence.
449 The data does not indicate a discernible long-term increase or decrease in global landslide impact;
450 rather the record shows that there is considerable inter-annual variability in global landslide event
451 incidence. The more active years have been associated with recognised regional patterns of rainfall, in
452 part driven by global climate anomalies, but there is no simple relationship with, for example, ENSO.
453 Relating climate modes to patterns of landsliding is challenging because of climate complexity and
454 change, requiring 30 year + datasets. Increased understanding of the impact of ENSO diversity on

455 regional climate, will improve models forecasting seasonal rainfall distribution and landslide impact.
456 This is particularly important in acutely affected areas such as India, China and Nepal.

457 Human disturbance (land use change) may be more detrimental to future landslide incidence
458 than climate change (Crozier, 2010; Anderson and Holcombe, 2013), and this is evidenced by a
459 number of studies (Innes, 1983; Glade, 2003; Soldati *et al.*, 2004; Imaizumi *et al.*, 2008; Borgatti and
460 Soldati, 2010; Lonigro *et al.*, 2015). A comprehensive review of climate-landslide studies by Gariano
461 and Guzzetti (2016) found the majority of papers (80%) showed a causal relationship between climate
462 change and landslides. However, the authors highlight the significant uncertainties surrounding our
463 current understanding of climate-landslide interaction. Specifically, the limited geographic scope of
464 research, challenges in downscaling climate scenarios to slope stability models, and complex
465 interactions between natural and human induced drivers of landslide activity. Gariano and Guzzetti
466 (2016) demonstrate that different climate variables will effect different landslide types and slope
467 settings. There is a high confidence that glacial retreat and permafrost degradation will increase slope
468 instabilities in high mountain areas in the long term, and high confidence that changes in heavy
469 precipitation will affect some regions. However, there is low confidence in projections for shallow
470 landslide activity in temperate and tropical regions, because of the coincident effects of human land
471 use practise (Seneviratne *et al.*, 2012). Further research is required to evaluate the impact of climate
472 change and human disturbance in different localities.

473 Our analyses have demonstrated that fatal landslide occurrence triggered by human activity is
474 increasing, driven by construction, illegal mining and illegal hillcutting. Fatal landslides occur when
475 construction and mining: (1) do not apply appropriate slope engineering, (2) mismanage spoil and, (3)
476 do not undertake a feasibility assessment (Hearn and Shakya, 2017). Appropriate building regulations
477 that account for the geo-environmental context of the settlement, provide clear guidance on
478 engineering and are enforced by local technical experts, are paramount in managing landslide risk
479 associated with urbanization and natural resource exploitation.

480 Holcombe et al. (2016) emphasised that planning policy alone is not sufficient to control
481 landslide risk in developing nations. This is due to the rapid and informal nature of construction, and
482 low-income of residents, who cannot finance expert guidance when building their homes. Settlements

483 are often built on hazardous land around urban centres and on roadsides, because of the benefits of
484 service access and employment opportunities (Smyth and Royale, 2000; Oven *et al.*, 2008; Lennartz,
485 2013; Anhorn *et al.*, 2015). Hillcutting is the dominant driver of instability during informal
486 construction (Holcombe *et al.*, 2016), and our results indicate fatal landslide events triggered by
487 hillcutting are increasing in Bangladesh, India and Nepal. Several landslides were triggered when
488 people cut slopes to collect coloured clay to decorate their houses for religious festivals. Here,
489 communication of landslide risk by local non-governmental organisations (NGOs) could prevent
490 future fatalities from this practice. Where governments are limited in capacity at a local level, NGOs
491 are important in implementing disaster risk reduction (Jones *et al.*, 2016), such as supporting
492 community-based slope engineering (*e.g.* Mosaic; Anderson and Holcombe, 2006).

493 Reporting of fatal landslides is likely to increase with the global growth in mobile technology
494 and internet access, particularly in remote mountain regions. Furthermore, advances in web mining
495 (data retrieval from the internet based on search criteria) and text mining (transforms unstructured
496 data into structured to discover knowledge) using machine learning offer methods to improve capture
497 of landslide reporting and data evaluation (*e.g.* Bhatia and Khalid, 2008; Kumar *et al.*, 2018). Global
498 landslide databases are designed to capture general trends in landslide occurrence rather than provide
499 data for local quantitative risk assessment. Continued collection of the database will develop our
500 understanding of the effect of climate and human disturbance on global landslide impact. The dataset
501 is a useful tool in identifying acutely landslide prone parts of the world and specific local drivers of
502 landslide impact; thereby highlighting locations which would benefit from further development in
503 early warning technology, landslide risk assessment and community capacity building. This is in
504 support of the future directions of the International Consortium on Landslides (Alcantara-Ayala *et al.*,
505 2017).

506

507 **Figures**

508 **Figure 1.** (a) The location of non-seismically triggered fatal landslide events from 2004 to 2016.
509 Individual landslide events shown by a black dot. (b) Number of non-seismically triggered fatal
510 landslide events from 2004 to 2016 by country. (c) The Gross National Income per capita (US \$) by
511 country (World Bank, 2018), and the location of major urban centres globally (ESRI, 2018).

512 **Figure 2.** The occurrence of non-seismically triggered landslide events from 2004 to 2016, and
513 cumulative total of recorded events. The data are arranged by pentads (five-day bins), starting on the
514 1st January each year, thus the first pentad includes records for 1-5 January, and there are a total 73
515 pentads. A simple 25-day moving average is shown.

516 **Figure 3.** (a) The occurrence of rainfall triggered landslide events from 2004 to 2016 (blue). The data
517 are arranged by pentads (five-day bins), starting on the 1st January each year. A simple 25-day
518 moving average is shown. The 25-day moving average for all non-seismically triggered landslide
519 events, is shown in black. (b) The occurrence of NSNR landslide events from 2004 to 2016 (purple).
520 The data are arranged by pentads (five-day bins), starting on the 1st January each year. A simple 25-
521 day moving average is shown. The 25-day moving average for all non-seismically triggered landslide
522 events, is shown in black.

523 **Figure 4.** Mean number of landslides per pentad through the annual cycle for all rainfall-triggered
524 landslides, and by geographical region. The 20th pentad is the 6-10 April, the 40th pentad is 15-19
525 July and 60th pentad is 23- 27 October.

526 **Figure 5.** Mean daily rainfall (mm) by month between 2004 and 2016 summarised by geographical
527 sub-region (blue bars). Global Precipitation Climatology Centre data (Xie *et al.*, 2013; GPCC, 2018)
528 was processed in ESRI ArcMap and Matlab. Mean daily rainfall-triggered landslide event occurrence
529 by month between 2004 and 2016 (black line). Daily values are used to overcome the difference in
530 month length.

531 **Figure 6.** (a) The occurrence of non-seismically triggered landslide events from 2004 to 2016: 25 day
532 and 1 year moving average (see also Fig. 2). (b) The number of fatalities from non-seismically

533 triggered landslide events from 2004 to 2016 by pentad with 25 day moving average. (c) Number of
534 single fatality landslides 2004 to 2016. (d) Number of landslide events incurring 64 to 128 fatalities
535 per event from 2004 to 2016. (e) Comparison of the complete landslide series (Fig. 6a) and multi-
536 fatality landslide series (excluding the 64 to 128 fatality class). (f) Anomalies in landslide event
537 occurrence by year by geographical region (multi-fatality events). (g) Anomalies in landslide
538 occurrence by year, by geographical region (single fatality events). Values greater than 1 standard
539 deviation from the mean are shown by a grey circle. Values greater than 2 standard deviations from
540 the mean are shown by a black circle.

541 **Figure 7.** Distribution of triggers of NSNR landslide events (770 events).

542 **Figure 8.** Number of landslide events triggered per year by (a) construction, (b) mining, (c) illegal
543 hill-cutting, (d) illegal mining, (e) legal mining and (f) mining (not specified). The black series
544 contains only multi-fatality landslide events. The grey series contains single and multi-fatality
545 landslide events.

546 **Figure 9.** By country, the number of landslide events triggered by (a) construction, (b) mining, (c)
547 illegal mining and (d) illegal hill-cutting, between 2004 to 2016.

548

549

550 **Data availability**

551 The GFLDe (2004 to 2016) is not currently available to the public, however a web-platform is under-
552 development to host the data openly at the University of Sheffield, UK. Due for release in 2018.

553

554 **Appendix A**

555 **Figure A1.** Sample autocorrelation plot for the pentad rainfall-triggered landslides. The 99%
556 confidence interval is shown by the blue horizontal lines.

557 **Figure A2.** Sample autocorrelation plot for the pentad NSNR landslides. The 99% confidence interval
558 is shown by the blue horizontal lines.

559 **Appendix B**

560 **Table B1.** Hierarchical linear regression results comparing the impact of seasonality in geographical
561 regions with the global mean number of landslides per pentad through the annual cycle (see Fig. 4).
562 The data series for each geographical region are sequentially added into the regression (such that the
563 second row of the table is a regression of S.Asia + S.E. Asia with the global series).

564

565 **Author contribution**

566 DP developed the methodology (2002-2003) and has consistently collected the database since 2004.
567 MF analysed the data and wrote up the results for this submission. DP contributed to writing.

568 **Competing interests**

569 The authors declare that they have no conflict of interest.

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573

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883

884 Tables

885 Table 1 Landslide trigger classification.

Classification	Definition	Keyword search terms
unknown	No trigger or obvious cause specified	-
rainfall	Rainfall raises pore-pressure in slope materials triggering failure.	"rain", "sleet", "storm", "hurricane", "precipitation", "flood", "water", "torrent"
earthquake	Strong ground motion associated with earthquakes weaken slope materials triggering failure (coseismic landslides).	"earthquake", "aftershock", "seismic", "tremor"
illegal mining	Unregulated or informal mining of slope materials in designated quarry or mine, where permission to extract material has not been granted.	"illegal", "permit", "regulat", "close", "informal", "pick", "illicit", "abandoned", "traditional", "license", "ban", "mine", "quarry", "spoil", "pit", "excavat"
illegal hillcutting	Hillcutting refers to the process of removing material from a hillslope for the purposes of altering its shape and/or to obtain slope material for use in construction, manufacture or farming. It is differentiated from mining, because it occurs on slopes that are not within a designated site of mining or quarrying, instead hillcutting typically occurs on individual slopes on steep agricultural land or on man-made slopes such as those along transport routes. Hillcutting differs from construction, because slope modification does not follow an engineering design to ensure slope stability. Hillcutting is assumed to be undertaken in an informal, unregulated manner (this is frequently noted in landslide reports)	"hillcut", "illegal", "permit", "regulat", "informal", "illicit", "traditional", "license", "ban", "excavat"
legal mining	Regulated and/or permitted mining of slope materials in designated quarry or mine, where permission to extract material has been granted and operations are managed.	"legal", "permit", "regulat", "pick", "license", "mine", "quarry", "spoil", "pit", "excavat"
mining (unknown)	Slope materials are extracted from a designated quarry or mine, but the report does not make it clear if the extraction is permitted or not.	"quarry", "mine", "spoil", "pit", "excavat"
construction	Permitted modification of a slope for the purposes of a construction project undertaken by professional labourers, following planning approval	"excavat", "construction", "site", "road", "build", "dig", "labour"
conflict and explosion	Landslide triggered by the detonation of an explosive device during military combat	"bomb", "mine", "soldier", "army", "explode", "explosion", "war", "conflict"
leaking pipe	Utility pipes carrying water that have been damaged and leak water onto a slope surface or within the hillslope, compromising its stability.	"pipe", "leak", "burst"
garbage collapse	Collapse of piles of municipal waste onto people, where stability of waste piles was disturbed by the passage of a person or persons	"waste", "trash", "rubbish", "garbage", "dump", "pick"
recreation	Triggered by passage of a person or persons walking/ climbing over a hillslope for recreation.	"climb", "mountain", "expedition", "ascent", "trek"
human action (unspecified)	Landslide report refers to a person or people present on a hillslope that collapses, without specifying the reason people occupied the slope or the landslide trigger	"people", "person", "men", "women", "children", "occup"
animal activity	Occupation of slope by animal triggering failure, either by weight and movement of animal on slope surface, or by burrowing within the slope subsurface.	"animal", "burrow", "tunnel"
fire	Naturally occurring or man-made fires, typically occurring in dry climates on vegetated terrain.	"fire"
natural dam or riverbank collapse	Collapse of a riverbank or natural dam without an apparent trigger, but likely caused by pore pressures building over time to a critical threshold in response to water levels. Material typically fails into a body of water, and often generates a flood wave.	"river", "bank", "dam", "earth", "flood", "wave", "collapse"
freezing	Heavy snowfall and expansion of water in hillslopes due to freezing, acting solely or together to destabilise the slope.	"snow", "extreme", "freeze", "ice", "cold"
freeze thaw (temperature change cold to hot), snowmelt	Failure of slope materials in response to temperature rise, including landslides triggered by the melting of snow or permafrost (in a non-volcanic setting)	"snow", "melt", "permafrost", "spring", "temperature"
volcanic eruption	Landslides (and mudflows) occurring in a volcanic environment triggered by volcanic activity- such as explosions and volcano-tectonic seismicity. This does not include events in active volcanic environments triggered by rainfall.	"volcan", "seismic", "activity", "eruption"
marine erosion	Triggered by sea erosion (only)- repeat wave impact	"coast", "sea", "erode"

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889 **Table 2.** Spearman’s rank correlation between mean daily rainfall and mean daily landslides by
 890 month (see Fig. 5).

Region	Correlation Coefficient	P value
Central America	0.8153	0.0012
South America	0.8062	0.0015
South East Asia	0.17	0.5974
South Asia	0.996	0
East Asia	0.9701	0

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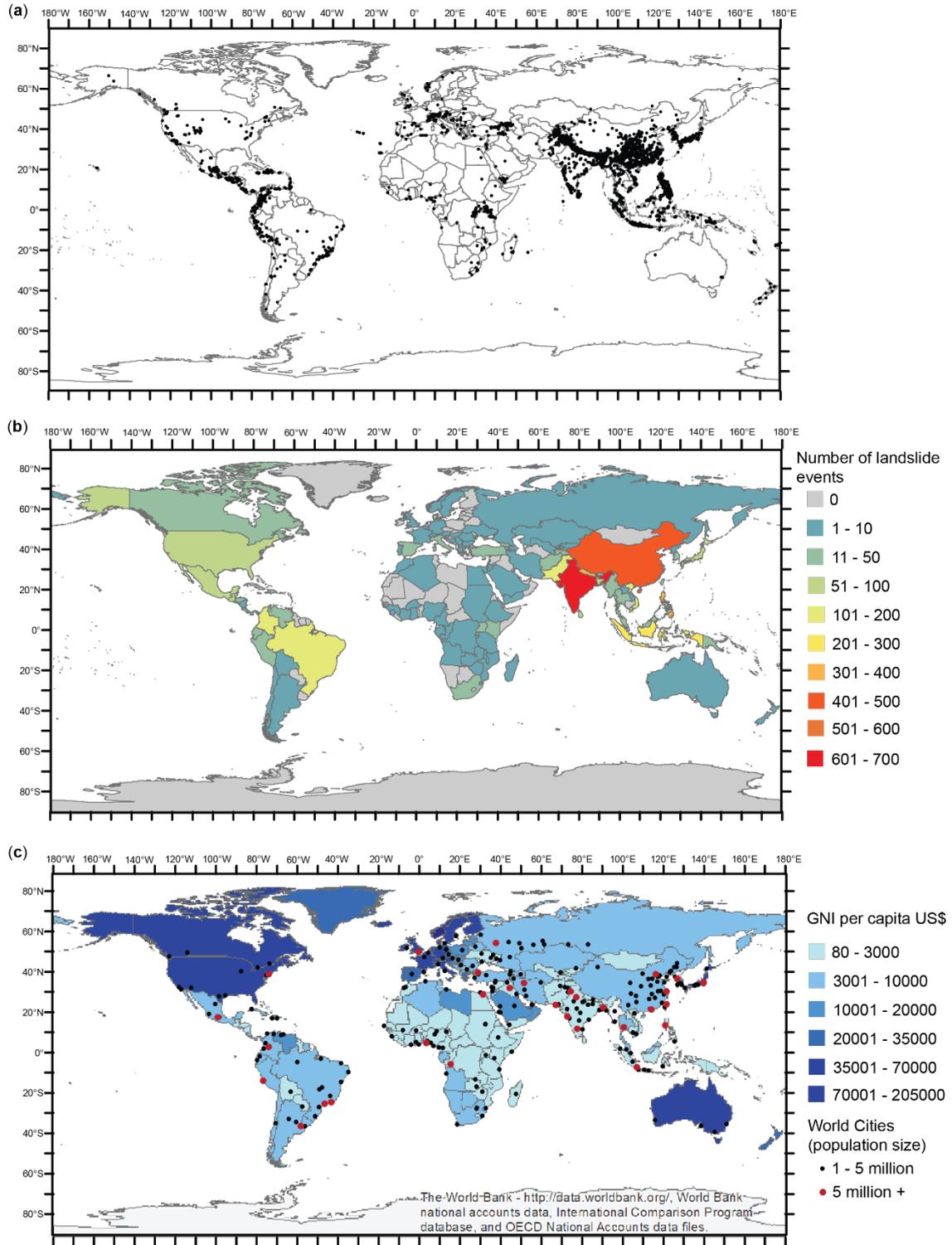
892 **Table B1.** Hierarchal linear regression results comparing the impact of seasonality in geographical
 893 regions with the global mean number of landslides per pentad through the annual cycle (see Fig. 4).
 894 The data series for each geographical region are sequentially added into the regression (such that the
 895 second row of the table is a regression of S.Asia + S.E. Asia with the global series).

Predictor Variables	N (cumulative)	% (of total N)	R ²	ΔR ²
+ S. Asia	1295	31.50	0.4962	
+ SE. Asia	2121	52.27	0.7365	0.2403
+ E. Asia	2804	71.88	0.8618	0.1253
+ S. America	3145	82.25	0.9129	0.0511
+ C. America	3340	88.03	0.9575	0.0446

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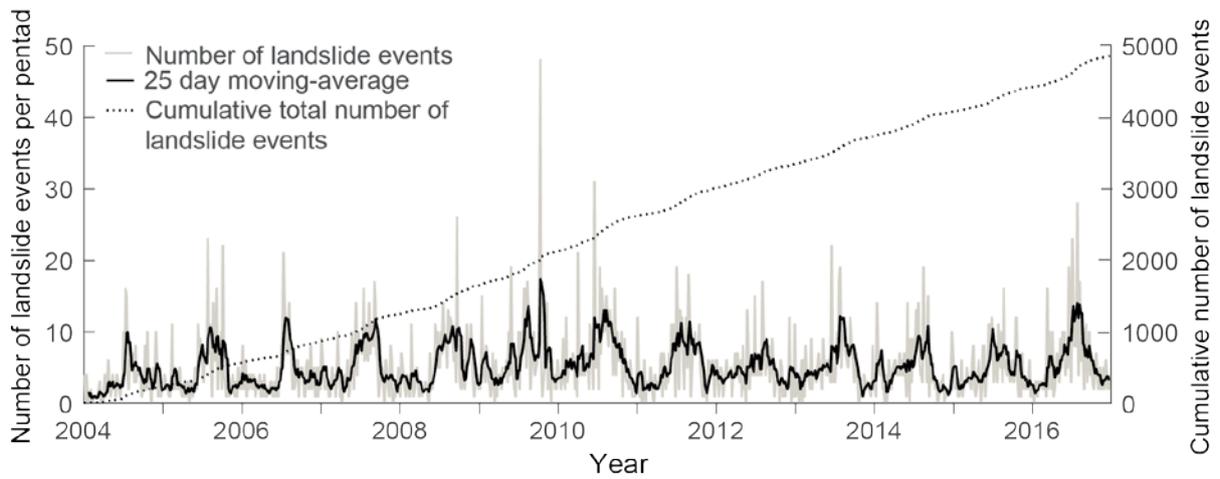
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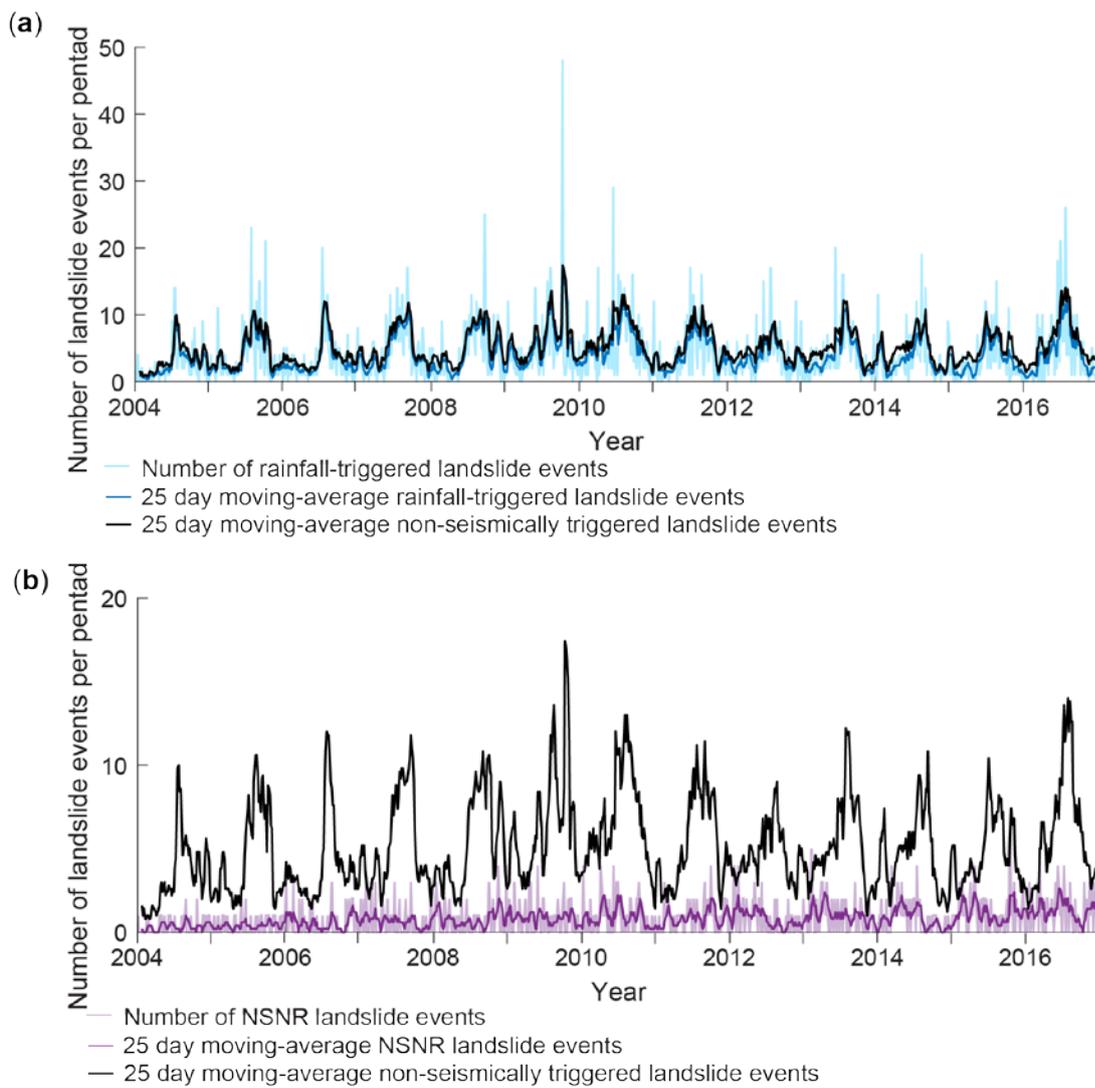
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902 **Fig. 2**



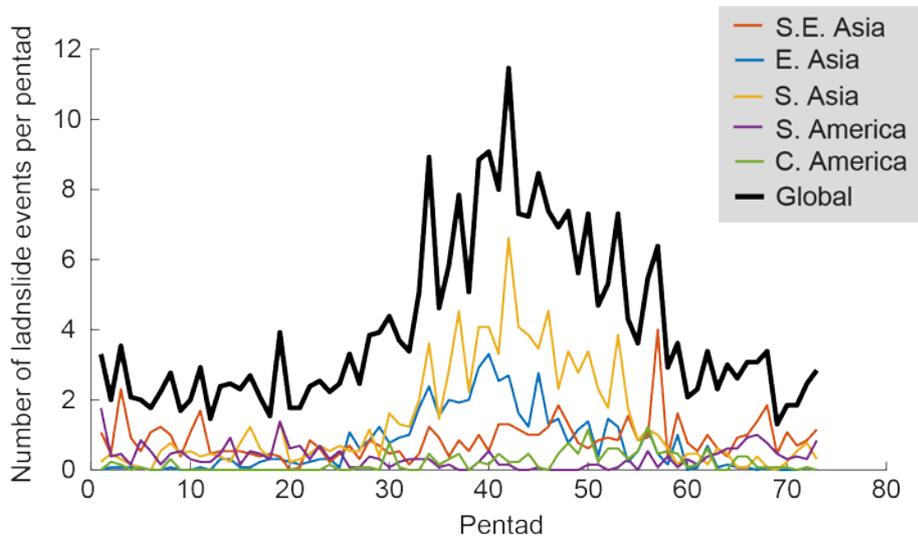
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904 **Fig. 3**



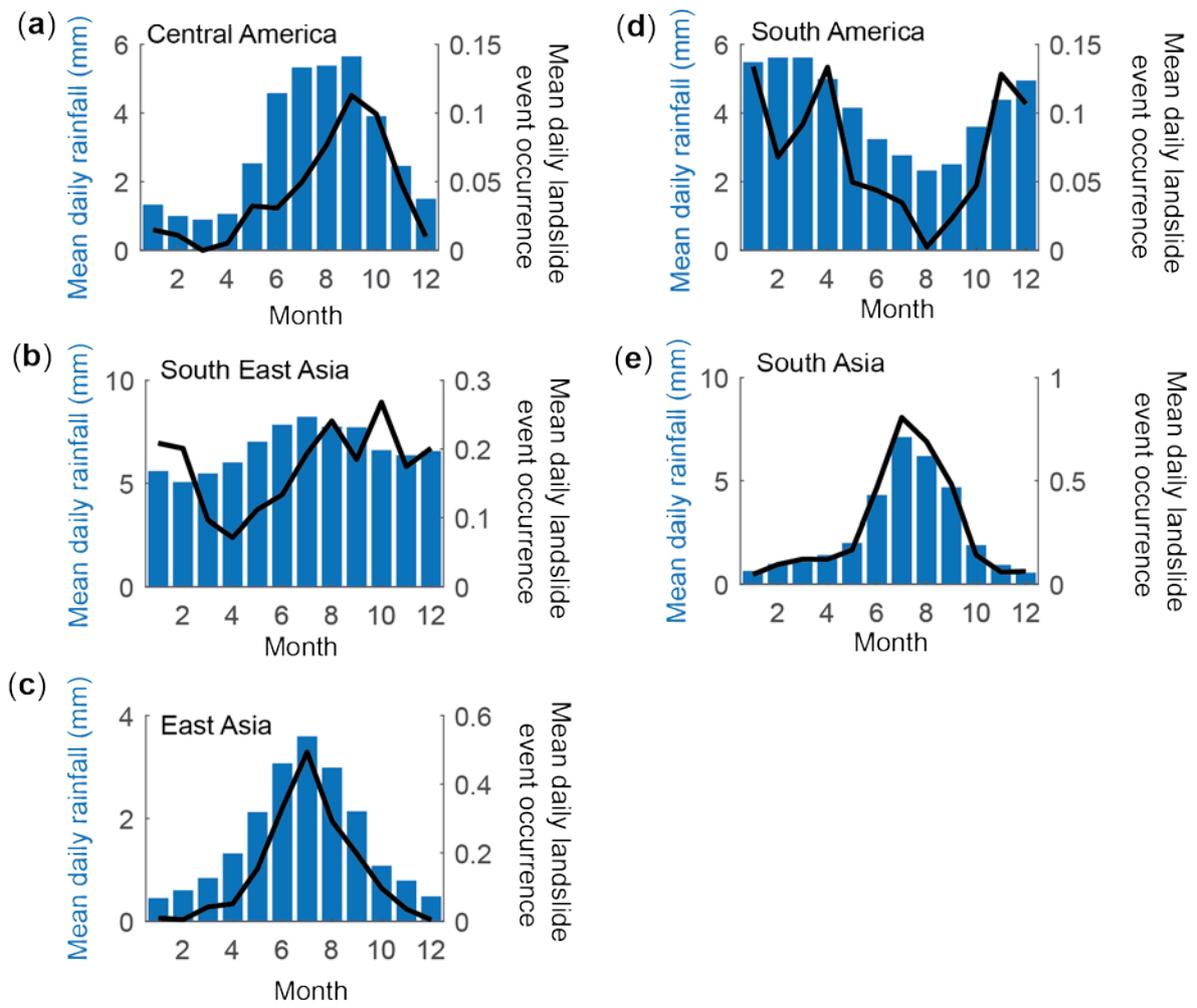
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906 **Fig. 4**

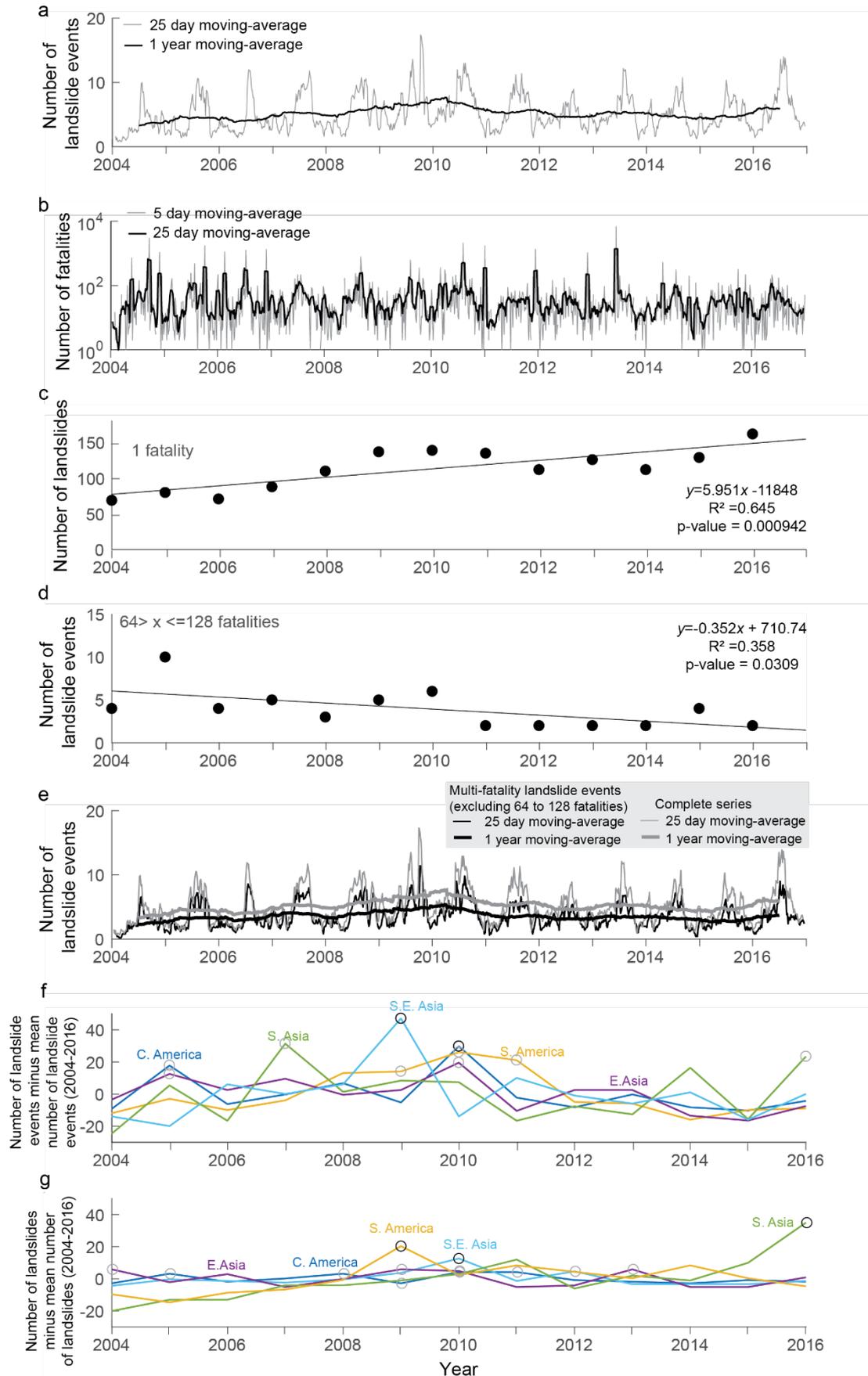


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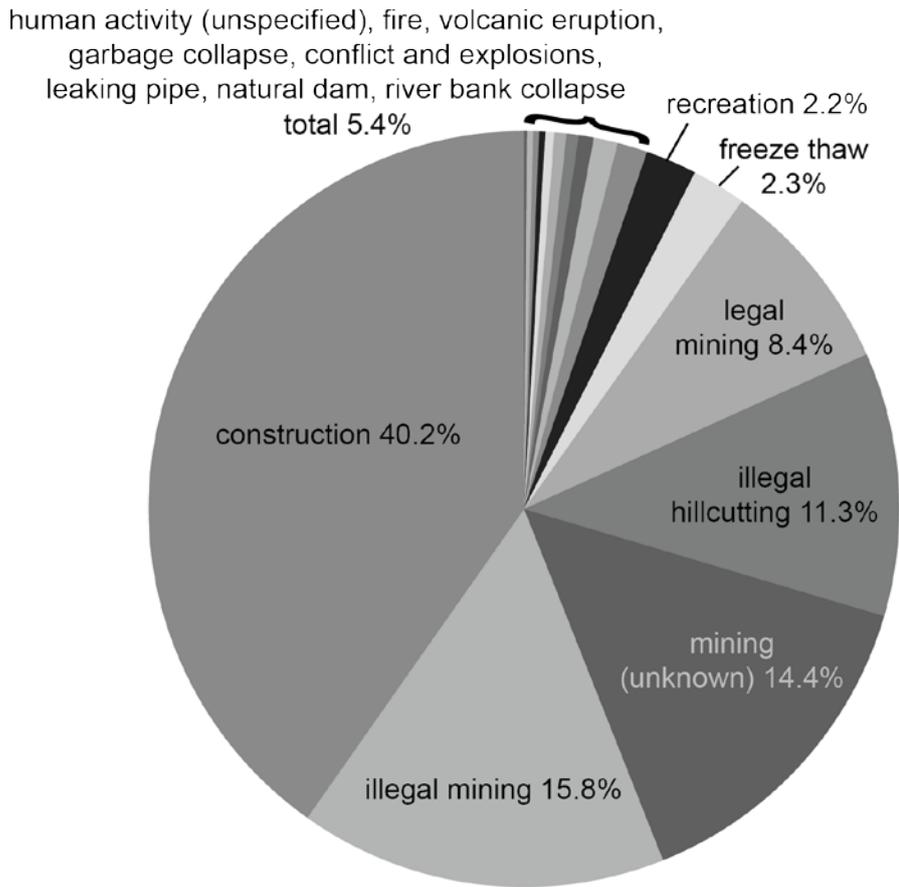
908 **Fig. 5**



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912 **Fig. 7**



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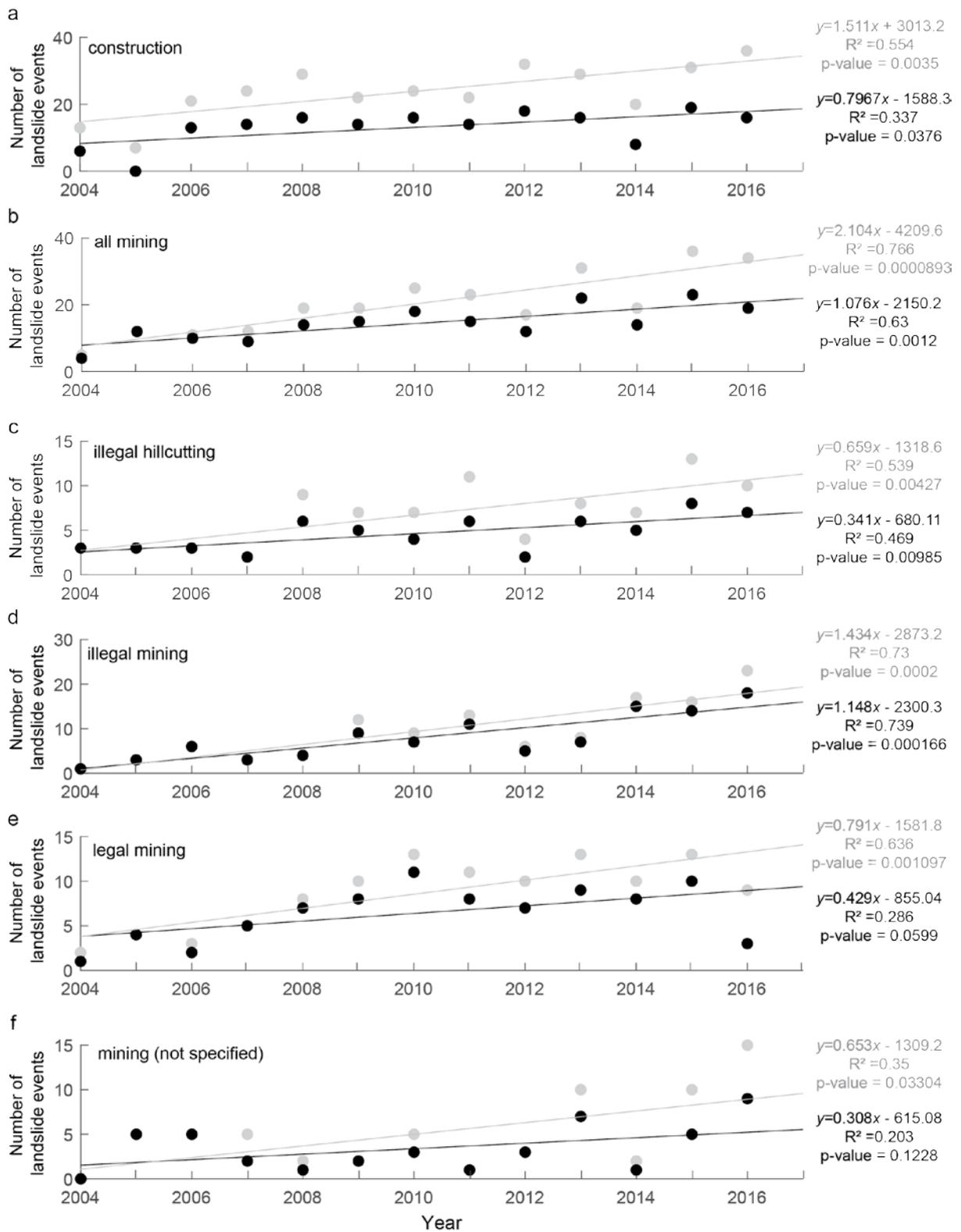
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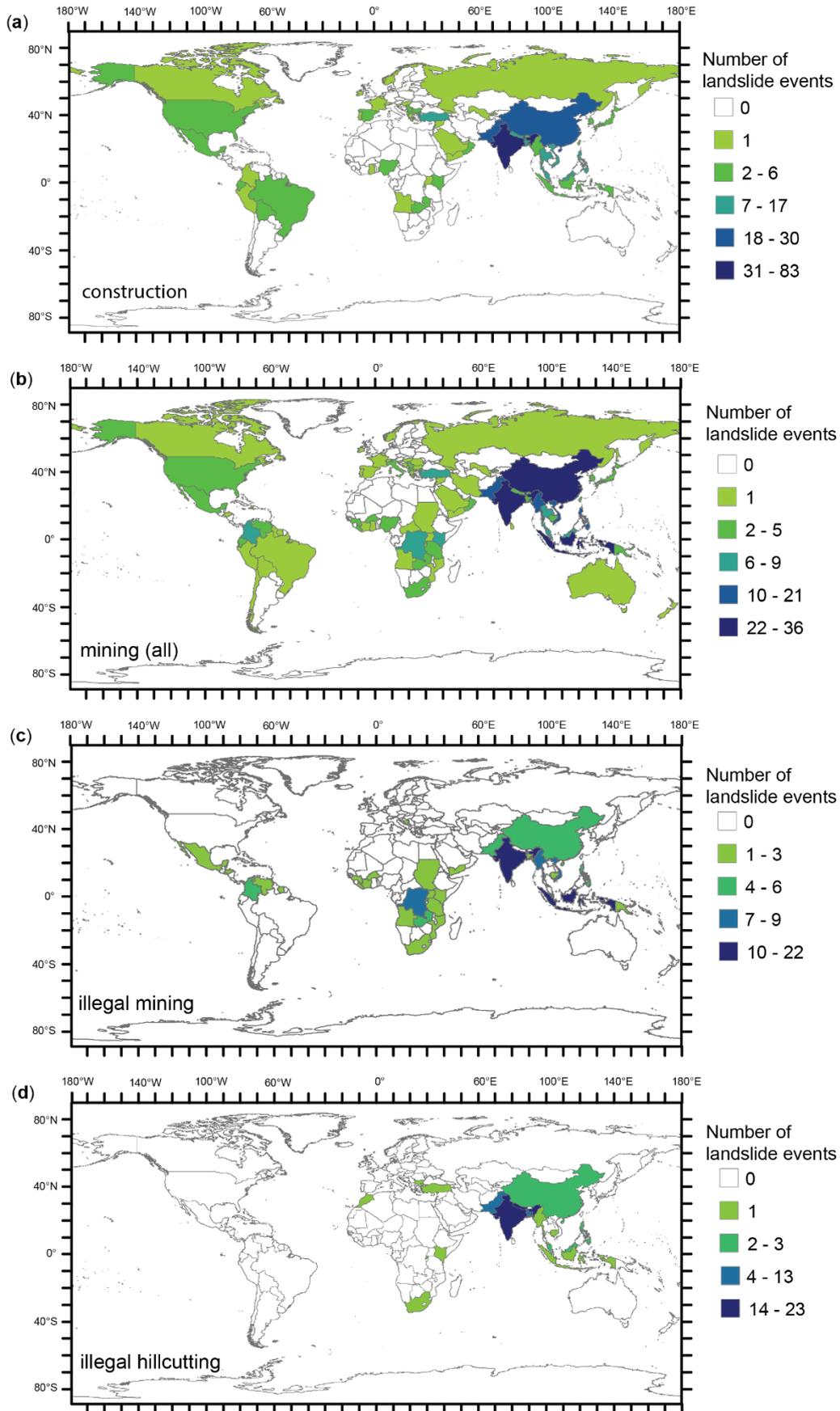
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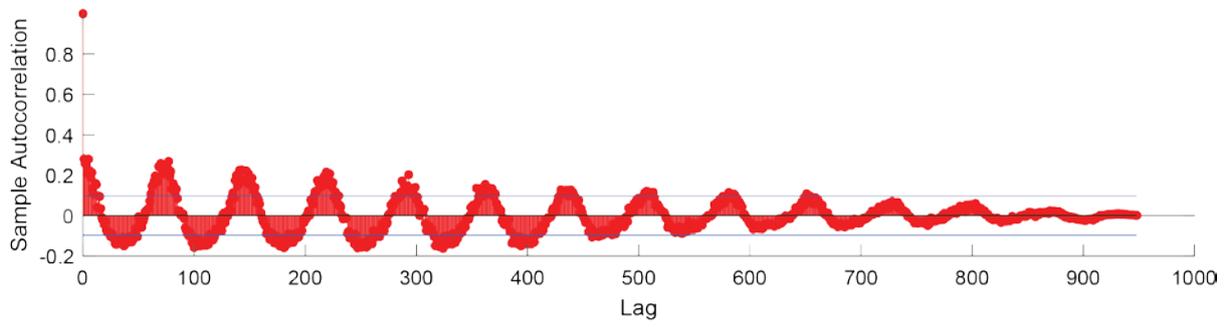
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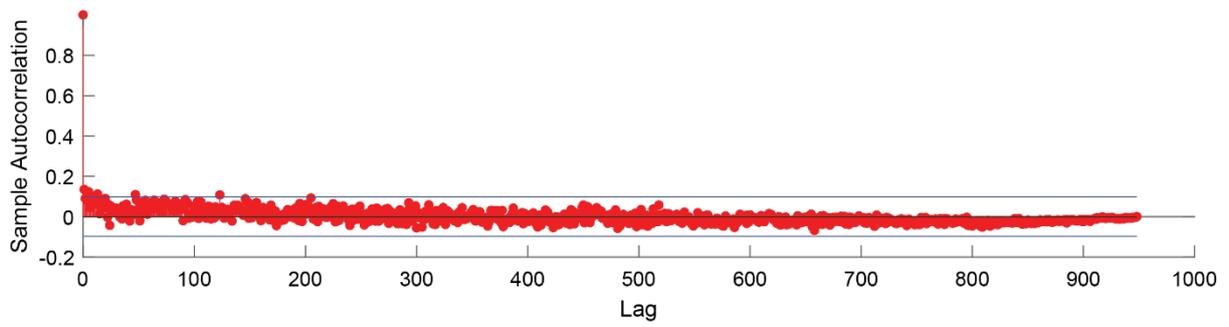
929 **Appendix Figures**

930 **Fig. A1**



931

932 **Fig. A2**



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