



1 **Influence of shearing rate on the residual strength characteristic of**
2 **three landslides soils in loess area**

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19 **Abstract**

20 In order to investigate the effect of the shearing rate on the residual shear strength of slip zone soils, a series ring shear tests
21 were carried out on slip zone soils from three landslides in loess area at the two shearing rates (0.1mm/min and 1 mm/min).
22 The slip zone soil specimens used in present study were from the northwest of China. Results indicated that the shear
23 displacement to achieve the residual stage for specimens with higher shearing rate is greater than that of the lower rate.
24 Relationship between the residual friction coefficients and normal stress shows that the residual friction coefficients for all
25 specimens under the lower normal stress were greater than that under the higher normal stress at two shearing rate.
26 Furthermore, the difference in the residual friction angle ϕ_r at the two shearing rates, $\phi_r(1) - \phi_r(0.1)$, under each normal
27 stress level were either positive or negative values, with the maximum absolute value of $\phi_r(1) - \phi_r(0.1)$ reach up to 2.218° .
28 However, the difference $\phi_r(1) - \phi_r(0.1)$ under all normal stresses was negative, which indicates that the residual shear
29 parameters reduced with the increasing of the shearing rate in loess area.

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33 **Keywords:** Residual shear strength; Ring shear test; Shearing rate; Normal stress; Slip zone soils



34 1. Introduction

35 Residual strength of soil is of great significance for evaluating the reactivating potential of the slope, in which consists of
36 pre-existing sliding surface. Residual strength of a landslide soil is defined as the minimum constant value of strength along
37 the slip plane, in which the soil particles are reoriented and subjected to sufficiently large displacements in relatively low
38 shearing rate (Skempton, 1985).

39 Numerical studies have been done to assess the residual strength through the laboratory tests using ring shear tests and
40 reversal direct shear tests (Chen and Liu, 2013; Vithana et al., 2012). It is a generally accepted fact that the measurement of
41 the residual strength is most preferred done with a ring shear test since it allows the soil specimen be sheared at unlimited
42 displacement which can simulate the field conditions more accurately (Lupini et al., 1981; Tiwari and Marui, 2005; Bhat,
43 2013; Sassa et al., 2004). Until now, several relationships between the residual strength and soil index parameters have been
44 reported in the literature with a wide range of soil by using various kinds of ring shear apparatus (Hoyos et al., 2014; Jiang et
45 al., 2016; Kimura et al., 2015; Li et al., 2013; Skempton, 1964). Furthermore, many studies have shown that the shearing rate
46 may or may not affect the minimum value of soil strength at residual states (Suzuki et al., 2007; Grelle and Guadagno,
47 2010; Gonghui et al., 2010; Bhat, 2013; Tika and Hutchinson, 1999; Lemos, 1985; Morgenstern and Hungr, 1984; Tika, 1999).

48 From the high shearing rate aspect in the geotechnical literatures, Morgenstern and Hungr (1984) carried out ring shear
49 tests on two types of coarse sand in high velocity and found that the frictional behavior was not affected by either the
50 velocity or the normal stress. However, there were many researchers asserted that the effect of shearing rate on the shear
51 behavior of soil cannot be ignored. For example, Skempton (1985), Tika and Hutchinson (1999) found that the faster
52 shearing rate above 100 mm/min may bring about great qualitative changes in the residual behavior. Moreover, Tika et al.
53 (1996) conducted fast ring shear tests on a wide range of natural soils and concluded that there are three types of rate effects
54 on the residual strength, namely, a positive rate effect (the residual strength of soil at fast rate is higher than that of the slow
55 rates), a neutral rate effect (the residual strength of soil is independent of the shearing rate) and a negative rate effect (the
56 residual strength of the soil at higher speed is lower than that of the lower speed). Recently, Gratchev Ivan and Sassa (2015)
57 reported that the residual strength of the clay decrease with the shear rate increase from 0.2 to 5 mm/s.

58 On the other hand, in the slow shearing rate range, Skempton (1985) reported that variation in the value of the residual



59 friction angle for shearing rates in a range of 0.05 to 0.35 mm/min was less than a 5% and concluded that the impact of
60 shearing rate on the residual strength of clay is almost negligible within slow rate displacement. Similarly, Bhat (2013)
61 concluded that there is hardly increase in residual strength of kaolin clay with the shearing rate ranging from 0.233 mm/min
62 to 0.586 mm/min. Furthermore, Yokota et al. (1995) showed that residual strength is not affected by shearing rate lower than
63 1.01 mm/min in ring-shear tests. Except the above studies, other similar results were also found in clays that the residual
64 strength is independent of the shearing rate (Chen and Liu, 2013; Tiwari and Marui, 2001). However, Suzuki et al. (2001) has
65 reported the shearing rate ranging from 0.02 to 2.0 mm/min significantly affected the residual strength of kaolin clay and
66 mud stone. Moreover, Gonghui et al. (2010) also has reported that the residual shear strength of the weathered serpentinite is
67 positively dependent on the shear rate in the slow rate.

68 On general, the effect of the shearing rate on the residual strength of the soil has not been sufficiently studied in high and
69 slow shearing rate range. Furthermore, except for a few studies, researchers have not widely reported the impact of the
70 shearing rate on the residual strength of loess soil in relatively lower shearing rate range from 0.1mm/min to 1 mm/min.
71 However, it should be noted that the residual strength parameters obtained from using different shearing rate may be adopted
72 to provide a guide for designing some precision engineering which require high accuracy of the design parameters, thus, the
73 effect of the shearing rate on the residual strength of soils should be fully understood to determine the parameters with high
74 reliability. In addition, residual strength of soil plays a key role in assessing the stability analysis and evaluating the
75 reactivation potential of landslides which consists of pre-existing slip plane surface. Therefore, accurate determination of the
76 residual strength parameters and their dependence on the shearing rate may affect the stability evaluation of landslides. Thus,
77 it is necessary to study the residual strength variation of loess in rate of shearing in order to have a good understanding of the
78 suitable approach for the residual strength measurement.

79 In this backdrop, the present study investigated the effect of the shearing rate on the residual strength of soil samples
80 obtained from three landslides in
81 loessic- developed areas at two different shearing rates (0.1mm/min and 1 mm/min) by using a ring shear apparatus. The
82 main objective of this study was to examine the change in the residual strength parameters of loess at different shearing rates
83 and their relationship with the normal stress in naturally drained ring shear tests.



84

85 **2. Geological setting of landslide sites**

86 Soil samples from three reactivated landslides in the northwest of China were selected for this study. Soil samples used
87 for the ring shear tests and index measuring tests predominantly consist of loess deposits and were collected in a disturbed
88 condition. For convenience, the names of landslide sites were abbreviated into Djg, Ydg, and Dbz. Figure 1 shows the study
89 sites and some views of the landslides.

90 **Dingjiagou landslide**

91 The Djg landslide, located at the mouth of Dingjia Gully in Yan'an of China, is geologically composed of upper loess
92 and lower sand shale in the Yan-chang formation. The dustpan-shaped landslide is inclined to the east, with its inclination
93 75.85° . The landslide is 350 m in width, 180 m in length, 70 m in elevation. The average thickness of slip mass is around 20
94 m, and the volume of landslide totaled approximately $105 \times 10^4 \text{ m}^3$. The slip mass is mainly constituted by loess, whereas the
95 sliding bed consists of sand shale in Yan-chang formation. The thickness of the sliding zone varied from 30 to 50 cm. The
96 front lateral region of the main slide section of the Djg landslide, where the sampling was performed, was found to be silty
97 clay.

98 **Yandonggou landslide**

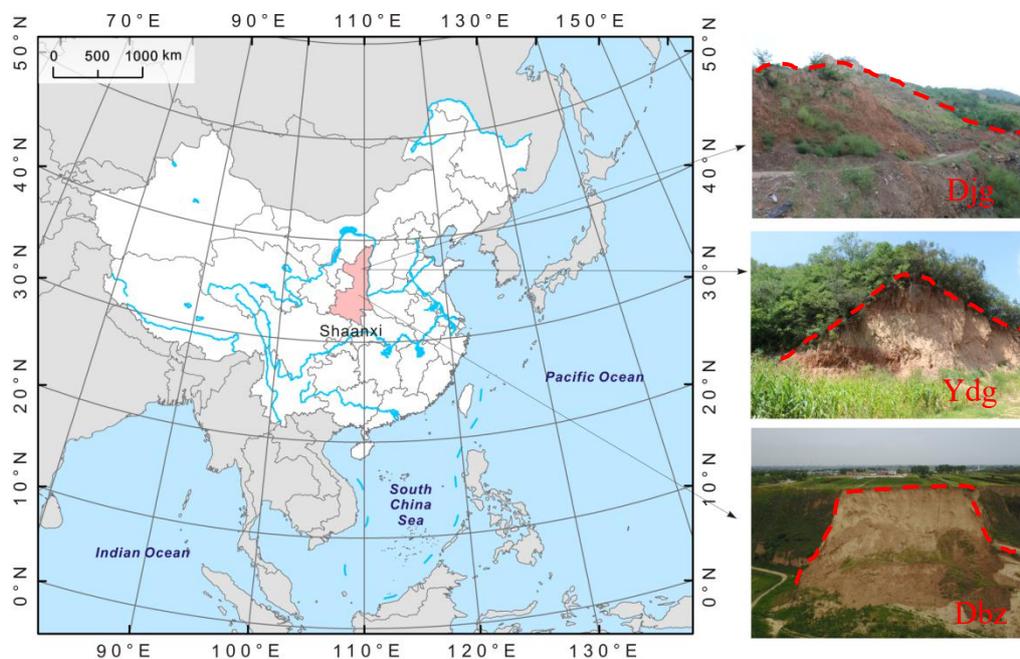
99 The Ydg landslide is located in the Qiaogou town of Yan'an in Shaan xi province of China. The top and the toe altitude of
100 the landslide are about 1165 m and 1110 m above the sea level, with the height difference between the toe and the top of
101 landslide about 55 m. The slides have well-developed boundaries with the main sliding direction of 240° and slope angle of
102 30° . From the landslides profile, the sliding masses from top to bottom were classified by Q_3 loess, Q_2 loess and clay soil,
103 respectively. Multiple landslide activities had occurred in this site, and the soil samples used in this study were collected
104 from Q_2 loess stratum within the slide ranged from 4.5 to 18 m in height.

105 **Dabuzi landslide**

106 The Dbz landslide is located in the middle part of Shaanxi province (about $108^\circ 51' 36''$ east longitude and $34^\circ 28' 48''$ north
107 latitude), China, which is a semi-arid zone dominated by loessic geology. In this region, the investigated site is classified as a
108 typical loess tableland with quaternary stratum. The sedimentary losses in this area are grey yellow, and the exposure stratum



109 in this area has been divided into two stratigraphic units, namely, the upper late Pleistocene (Q_3) loess and the lower
110 mid-Pleistocene (Q_2) loess, of which the Q_3 loess is younger. The Q_3 loess is closest to the surface and is up to approximately
111 12 m thick, while the thickness of Q_2 loess may reach an upper limit of about 50 m (Leng et al., 2018). The loess in this area
112 have well-developed vertical joints (Sun et al., 2009). The travel distance and the maximum width of the slip mass are
113 roughly estimated to be 121.55 m and 133.46 m, respectively. The armchair-shaped landslide shows an apparent sliding
114 plane, with an area of approximately 15,660 m² and about 66.25 m maximum difference in elevation. The main direction of
115 this landslide is approximately 355°. The exposed slip zone in the side scarp of the landslide, where the sampling was done,
116 was found to be entirely in the Q_2 loess stratum of the Dbz landslide site. The thickness of narrow- band slip zone loess is
117 less than 1.0 cm, inclined at around 65° to the horizontal direction. Since the band of slip zone is thin, mix soils which consist
118 of the slip zone soils, the very thin upper and lower parts of the loess are mingled together and served as the representative
119 samples of the slip zone loess in this site.



120
121 Figure 1. Location of study sites and some views of landslides

122 3. Experimental scheme

123 3.1. Testing sample



124 The fact that the residual shear strength is independent of the stress history was reported by many researchers (Bishop et
 125 al., 1971;Vithana et al., 2012;Stark Timothy et al., 2005). Thus, disturbed loess samples from each of the three landslides
 126 weighing about 25 kg were collected from the slip surface soil of each slide and used to determine residual shear strength.

127 The soil samples were air-dried and then crushed with a mortar and pestle, and subsequently processing it through 0.5 mm
 128 sieve. Distilled water was added to the soil samples until desired density and water content were obtained. The physical
 129 parameters such as natural moisture content, specific gravity, bulk density, plastic limit, and liquid limit were determined in
 130 accordance with the Chinese National Standards (CNS) GB/T50123-1999 (standards for soil test methods) (SAC, 1999),
 131 but clay size was defined to be less than $2\mu\text{m}$ followed ASTM, D422 (ASTM, 2007). Each soil sample was separated into
 132 clay (sub 0.002 mm), silt (0.002-0.075 mm), and sand (0.075-0.5 mm) fractions. The physical indexes of the soil are listed in
 133 Table 1.

134 The grain size distribution of soil was measured using a laser particle size analyzer Battersize 2000 (Dandong Battersize
 135 Instruments Corporation, Dandong, China). The sieved soil samples were used to determine particle size distribution. In this
 136 study, soil samples were treated with sodium hexaphosphate, serving as a dispersant, to disaggregate the bond between the
 137 particles. The results show that the clay fraction in Djg landslide soil (24%) is more than two times than that from Ydg (9%)
 138 and Dbz (9.1%). Furthermore, the particle size analyses illustrates that the percentage of silt-sized soil in three landslides
 139 ranged from 75.66% to 87.4%. In addition, Ydg landslide soil consists of the greatest percentage of the sand fraction which
 140 reaches up to 10.55%.

141 In present study, a total of twenty four specimens were tested at two shearing rate (0.1mm/min and 1 mm/min) and under
 142 normal stresses ranged from 100kN/m^2 to 400kN/m^2 in a ring shear apparatus.

143 **Table 1** Physical parameters of slip-zone loess

<i>sites</i>	ρ_d	W	ρ	G_s	W_L	W_p	Grain size fractions (%)			
							<0.002m	0.002-0.005m	0.005-0.075	0.075-0.5m
Djg	1.74	19.	2.0	2.6	3	20	24	11.48	64.18	0.34
Ydg	1.47	18	1.7	2.7	3	19	9	5.28	75.17	10.55



Dbz	1.48	16	1.7	2.7	3	21	9.1	6.4	81	3.5
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144 Notes: ρ_d = dry density; w=moisture water content; ρ = bulk density; G_s = specific gravity; W_L =liquid limit; W_p = plastic limit

145

146 3.2. Testing apparatus

147 The advantage of a ring shear test apparatus to measure residual shear strength including its ability to allow
148 unidirectional shearing of a soil specimen (Bishop et al., 1971;Tika, 1999;Suzuki et al., 2007;Bromhead, 1979). Thus, a ring
149 shear apparatus was used in this study.

150 An advanced ring shearing apparatus (SRS-150) manufactured by GCTS (Arizona, USA) was adopted in ring shear
151 tests and the photos of apparatus were shown in Fig.2, which consists mainly of a shear box with an outer diameter of 150
152 mm, an inter diameter of 100 mm and the maximal sample height of 250 mm. The shearing box consists of the upper
153 shearing box and the lower shearing box. In the shearing process, the upper shearing box keeps still while the lower one
154 rotates. The apparatus, which provides effective specimen area of 98 cm², is capable of shearing the specimen for large
155 displacement in single direction. The annular specimen is confined by inside and outside metal rings. Moreover, the
156 specimen is confined by bottom annular porous plates and top annular porous plates in which have sharp-edged radial metal
157 fins which protrude vertically into the top and bottom of the specimen at the shearing process. The normal stress, shearing
158 strength and shearing displacement can be monitored while shearing by computer. The measurement features of the ring
159 shear apparatus employed in this study are described as follows: shearing rate range from 0.001 to 360 degrees per minute,
160 10 kN axial load capacity, 300 N.m continuous torque capacity, maximum normal stress of 1000 kN/m².

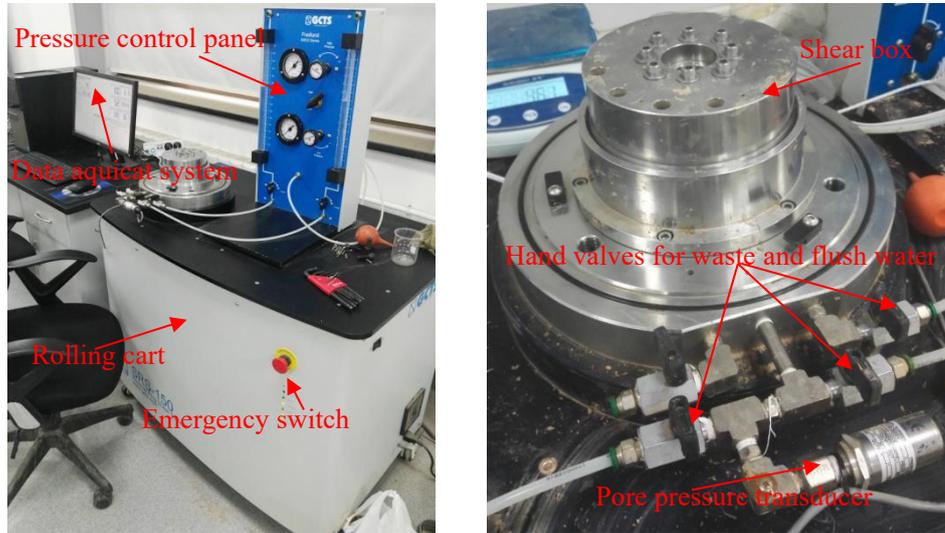


Figure 2. Ring shear apparatus (SRS-150)

3.3. Testing procedure

In present study, reconstituted samples of the sub 0.5 mm soil fraction were used in the testing as it was reported that the residual strength of the soil was unaffected by its initial structure (Bishop et al., 1971; Vithana et al., 2012). Specimens were first prepared by adding distilled water to the air-dried soil until the saturated moisture contents of the three landslide soils were obtained. Then, specimens were kept in a sealed container for at least one week to fully hydrate. Specimens are then reconstituted in the ring-shaped chamber of the apparatus by compaction. In order to make the sample uniform while packing, the sample was placed in three layers, and each layer was tamped under a vertical stress which is lower than the given normal stress to achieve the design height. The final height of the specimen in the ring shear apparatus after tamping varied but was typically about 20 mm (to achieve a specific bulk density). The specimen was then consolidated under a specific effective normal stress in a range of 100kN/m² to 400kN/m² until required consolidation was achieved. Then, the consolidated specimen is subjected to shearing under constant normal stress by rotating the lower half of the shear box attached to a gear, while the upper half remains still. In ring shear tests, the normal stress at the shearing was the same as at consolidation stage.

In this study, ring shear tests were performed in a single stage under drained condition and the samples were subjected to shear until the residual state was achieved. Drain condition of the shearing process is provided by two porous stones



178 attached on the top and the bottom platen of the specimen container. As for soil specimens with low permeability, the rate of
179 excess pore pressure generation in the shear box may exceeded that of pore-pressure dissipation, this type of condition is
180 identified as naturally drained condition in previous studies(Okada et al., 2004). Furthermore, Tiwari (2000) asserted that it
181 was acceptable to use a shearing rate below 1.1 mm/min to simulate the field naturally drained condition. Thus, shearing
182 rates of 0.1 mm/min and 1 mm/min were used in this study to simulate the naturally drained condition of the slip zone soils.

183 **4. Results and discussions**

184 Twenty four specimens were tested to investigate the shear characteristics of the slip-zone soils in the ring shear
185 apparatus. Tests results are shown in this section.

186 **4.1. Shear behavior**

187 Figures 3a, 4a and 5a show the typical shear characteristics of the slip-zone soils (shearing rate 0.1 mm/min and 1
188 mm/min) obtained from three different locations, where, the shear stress is plotted against the shear displacement at the
189 normal stress ranged from 100kN/m² to 400kN/m². It is a widely accepted fact that normal stress has effect on the shear
190 behavior of the soil (Kimura et al., 2015;Stark Timothy et al., 2005;Stark et al., 2005;Eid, 2014), thus, the shear behavior of
191 samples at the peak and residual stages, where, the determined peak friction coefficient as well as residual friction coefficient
192 are plotted in Figure 3b, 4b, and 5b against the corresponding effective normal stresses. The friction coefficient is defined as
193 the shear stress divided by the effective normal stress.

194 Figures 3a, 4a and 5a demonstrate that shear stress increases dramatically within small shear displacement and then
195 reduces with shearing displacement, until residual conditions were achieved at large displacements. Furthermore, it is clear
196 that peak strength as well as residual strength of the samples with high shearing rate is almost smaller than that of the
197 samples with low rate. In Figures 3a, 4a and 5a, a clear drop can be seen, at any normal stress, for specimens obtained from
198 all sites. It is obvious that Djg specimens showed greater peak-post drop than that of Ydg and Dbz specimens. According to
199 the conclusion that the residual stage is attained if a constant shear stress is measured for more than half an hour (Bromhead,
200 1992), it can be seen that the shear displacement to achieve the residual stage for specimens with higher shearing rate is
201 greater than that of the lower rate. For example, the minimum shear displacements for attaining residual condition for Djg
202 specimens with low and high shearing rate were about 360mm and 650mm, respectively. Under the shearing rate of



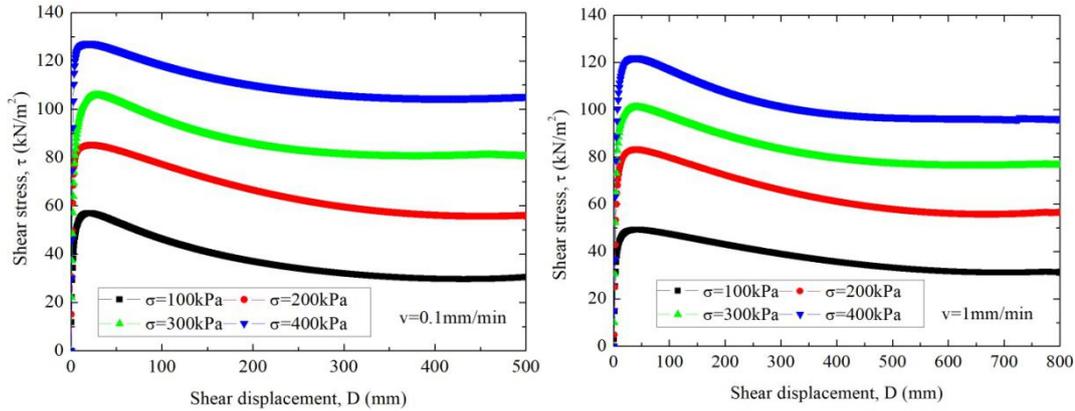
203 0.1mm/min and 1mm/min, Ydg specimens need approximately 80mm and 1,400 mm displacement to achieve residual stage.
204 However, Dbz specimens require about 40mm and 60mm displacement to reach residual condition for low and high shearing
205 rate, respectively.

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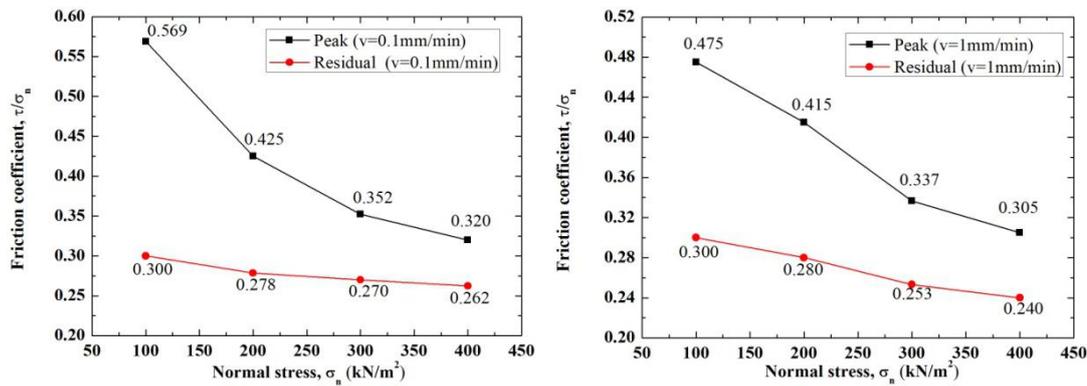
207 **4.2. Effect of normal stress on the friction coefficients**

208 It can be seen from the Figures 3b, 4b and 5b that the friction coefficients (peak and residual) are higher at lower
209 effective normal stress levels. For example, the peak and residual friction coefficient of Djg landslide soils at the shearing
210 rate of 0.1mm/min reduced from 0.569 to 0.32 and from 0.3 to 0.262, respectively. Similarly, results obtained from other two
211 landslides soils also showed that the friction coefficients decrease nonlinearly with the normal stresses. Furthermore,
212 specimens with shearing rate of 0.1mm/min attained greater friction coefficients than that with shearing rate of 1mm/min.

213 In order to get an insight into the effects of the normal stress on the slip zone shear strength, the shear behavior of the
214 soil sheared at the normal stress of 100kN/m² and 400kN/m² were selected for analysis. At the normal stress of 100kN/m²,
215 Djg samples showed about 47.3% and 36.8% decrease in the friction coefficient from the peak friction coefficient at the
216 shearing rate of 0.1 and 1 mm/min, respectively, which is greater than in the Ydg (about 9.8% and 10.3%) and Dbz (about
217 2.4% and 3.2%) samples. In Figures 3b, 4b and 5b, on average, it is obvious that the decrease of the friction coefficient from
218 the peak strength to the residual strength in the Djg sample was almost 18.1% and 21.3% for the sample consolidated at
219 normal stress of 400kN/m² under the shearing rate of 0.1mm/min and 1mm/min (Figure 3b), While the friction coefficient
220 reduction in Ydg sample with low and high shearing rate were only about 4.1% and 4.8% (Figure 4b). And the friction
221 coefficient reduction in Dbz samples with low and high rate were only approximately 5.6% and 6.0% (Figure 5b) from the
222 peak strength, respectively. Based on the conclusion that the post-peak drop in strength of soil is only due to particle
223 reorientation after the peak strength (Mesri and Shahien, 2003;Skepmtton, 1964), the results in this study demonstrated that
224 the Djg landslide soil existed the greater particle reorientation compared with that of other two landslide soils.

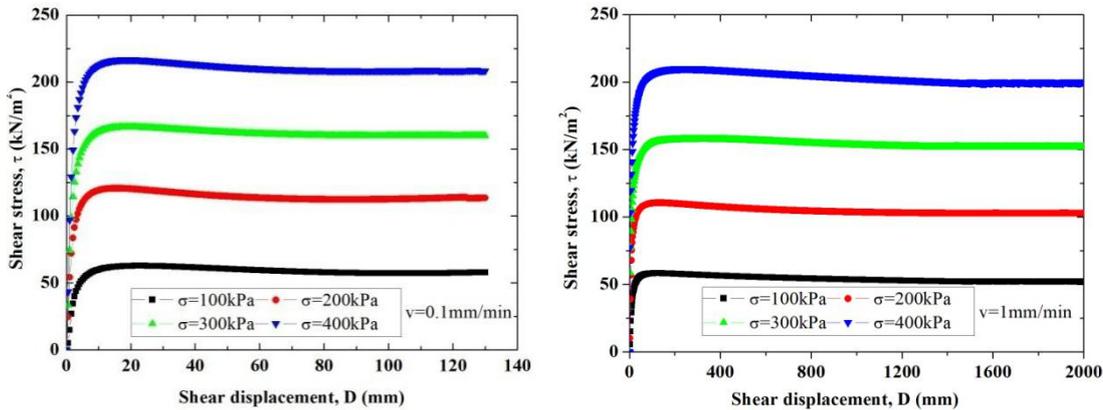


(a) Relationship between shear stress and shear displacement

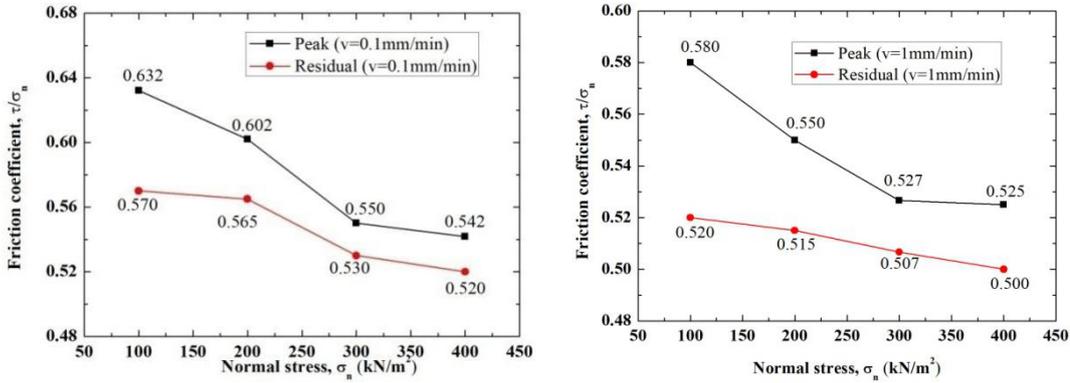


(b) Relationship between friction coefficient and normal stress

Figure 3. Shear behavior characteristics of Djg soil samples

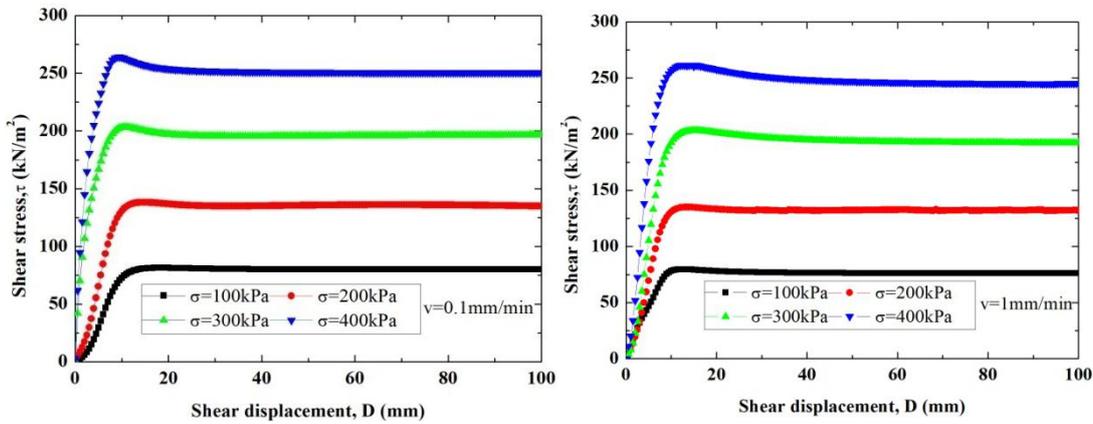


(a) Relationship between shear stress and shear displacement

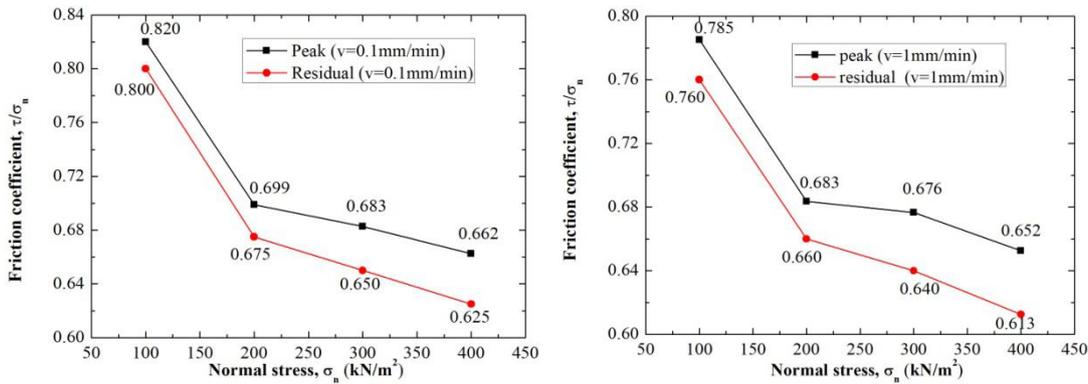


233
 234 (b) Relationship between friction coefficient and normal stress

235 Figure 4. Shear behavior characteristics of Ydg soil samples



236
 237 (a) Relationship between shear stress and shear displacement



238
 239 (b) Relationship between friction coefficient and normal stress

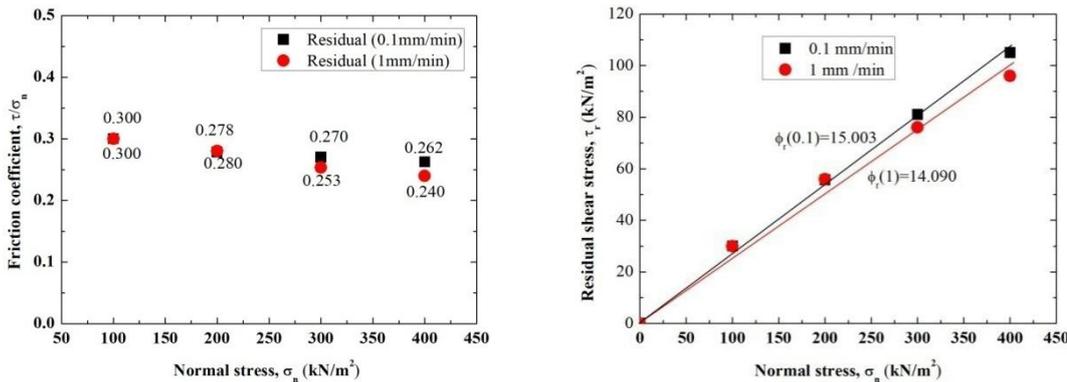
240 Figure 5. Shear behavior characteristics of the Dbz soil samples



241 **4.3. Effects of shearing rate on residual strength parameter**

242 For the representative samples described above, Figures 6, 7 and 8 show the relationships between the residual friction
 243 coefficient and the normal stress, and the residual strength parameters. The residual friction coefficient is plotted against the
 244 normal stress. The residual friction coefficient is defined as the residual shear strength divided by normal stress. It is widely
 245 recognized that the shear strength parameters including cohesion and friction angle (Terzaghi, 1951; Stark et al., 2005).
 246 However, according to the previous studies, the residual angle of soils varies depended on the soil properties as well as the
 247 magnitude of normal stress provided the residual cohesion of soil is zero (Skepmton, 1964; Bishop, 1971; Kimura et al., 2014).
 248 Thus, in this study, the residual frictions are calculated by Coulomb's law assumed the residual cohesion is zero. The
 249 residual strength parameters were defined as $\phi_r(0.1)$ and $\phi_r(1)$ at the low shearing rate and high shearing rate, respectively.
 250 And the difference between the residual friction angles at two shearing rate was defined as $\phi_r(1) - \phi_r(0.1)$. Comparatively,
 251 the residual friction coefficient was defined as $\tau_r/\sigma_n(0.1)$ at the low shearing rate and $\tau_r/\sigma_n(1)$ at the high shearing rate,
 252 respectively. Furthermore, the difference between the residual friction coefficients was defined as $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$. Table
 253 2 summarized the residual parameters of the landslide soils.

254 Figure 6 shows that the residual friction coefficients were relatively low in Djg samples. The coefficients $\tau_r/\sigma_n(0.1)$
 255 and $\tau_r/\sigma_n(1)$ at the normal stress of 100kN/m^2 to 400kN/m^2 ranged from 0.3 to 0.262 and from 0.3 to 0.24, respectively. The
 256 difference between the friction coefficients, $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$, at each normal stress level are varied in a range of -0.022 to
 257 $+0.002$. For the difference between the residual friction angles, $\phi_r(1) - \phi_r(0.1)$, ranged from -1.212° to $+0.079^\circ$ (Table 2). For
 258 normal stress above 200kN/m^2 , the coefficient $\tau_r/\sigma_n(0.1)$ was found to be greater in the magnitude than the coefficient $\tau_r/\sigma_n(1)$.

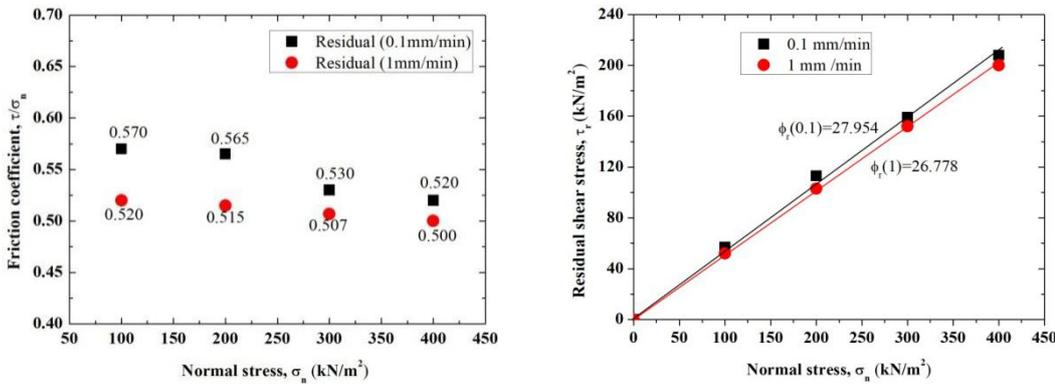


259



260 Figure 6. Relationships between residual shear stress and normal stress, and residual strength parameter for Djg soil
 261 sample

262
 263 Figure 7 gives the relationship between the residual coefficient and normal stress, and residual shear strength parameter
 264 for Ydg samples. The coefficients $\tau_r/\sigma_n(0.1)$ and $\tau_r/\sigma_n(1)$ under the normal stress of 100kN/m² to 400kN/m² ranged from 0.57
 265 to 0.52 and from 0.52 to 0.50, respectively. Furthermore, the difference $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$ at each normal stress was from
 266 -0.05 to -0.02. As for the difference between the residual friction angles, $\phi_r(1) - \phi_r(0.1)$, was in a range of -2.218° to -0.909°.
 267 In case of Ydg soil sample, there was insignificant reduction in residual friction coefficients with the increasing of shearing
 268 rate for all normal stresses.



269
 270 Figure 7. Relationships between residual shear stress and normal stress, and residual strength parameter for Ydg soil sample

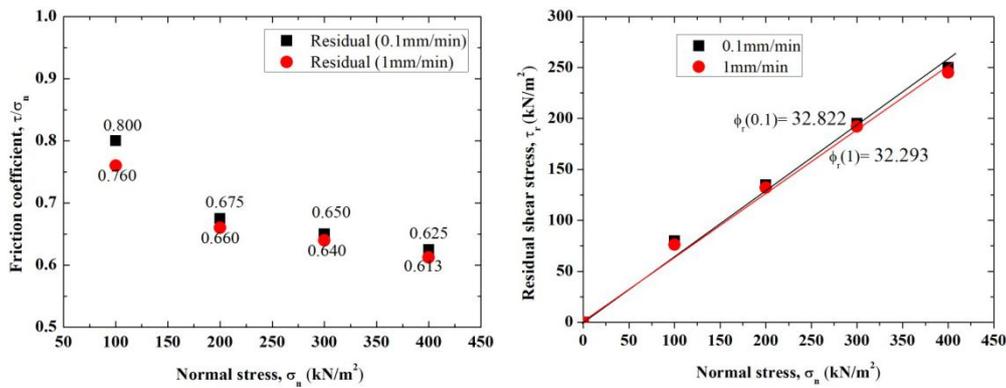
271
 272 Figure 8 presents the results of the Dbz samples. The coefficients $\tau_r/\sigma_n(0.1)$ and $\tau_r/\sigma_n(1)$ at the normal stress of 100kN/m² to
 273 400kN/m² ranged from 0.8 to 0.625 and from 0.76 to 0.613, respectively. The difference $\tau_r/\sigma_n(1) - \tau_r/\sigma_n(0.1)$ at each normal
 274 stress was from -0.04 to -0.01. The difference $\phi_r(1) - \phi_r(0.1)$ was from -1.425° to -0.405°. For Dbz samples, there was
 275 somewhat decrease of the residual friction coefficients with the increasing of the shearing rate for all normal stress levels. It
 276 is noted that the maximum difference was found at the lowest normal stress of 100kN/m².

277 Table 2 summarizes residual strength parameters including $\phi_r(0.1)$ and $\phi_r(1)$ of all specimens obtained from the ring
 278 shear tests in this study. As for the Djg samples, the residual strength parameter $\phi_r(0.1)$ and $\phi_r(1)$ for all normal stress were



279 found to be 15.003° and 14.09° , respectively. However, the residual friction angles $\phi_r(0.1)$ and $\phi_r(1)$ of the Ydg samples
 280 were obtained to be 27.954° and 26.778° , respectively. In the case of Dbz sample, the friction angles $\phi_r(0.1)$ and $\phi_r(1)$ were
 281 high, 32.822° and 32.293° , respectively. The residual friction angles $\phi_r(0.1)$ and $\phi_r(1)$ under all normal stresses were from
 282 15.003° to 32.822° and from 14.09° to 32.293° , respectively.

283 Due to the influence of the shearing rate, the difference $\phi_r(1) - \phi_r(0.1)$ in the Djg, Ydg and Dbz samples, were -0.913° ,
 284 -1.176° and -0.529° , respectively. Wang (2014) and Fan et al. (2017) asserted that the residual shear strength of remoulded
 285 loess hardly affected by shearing rate below 5mm/min. However, the results in this study shown that $\phi_r(1) - \phi_r(0.1)$ under
 286 all normal stress levels were negative for slip zone loess. Moreover, the maximum value of the difference $\phi_r(1) - \phi_r(0.1)$ even
 287 reached about 1.176° .



288
 289 Figure 8. Relationships between residual shear stress and normal stress, and residual strength parameter for Dbz soil sample
 290

291 Table 2 Residual shear strength parameter of landslide soils

No	Sample	Normal stress(kN/m ²)	Residual strength parameter				Difference in parameter	
			0.1 mm/min $\phi_r(0.1)$ ($c_r(0.1)=0$) (Degrees)		1 mm/min $\phi_r(1)$ ($c_r(1)=0$) (Degrees)		$\phi_r(1) - \phi_r(0.1)$ (Degrees)	
			Under each σ_n	Under all σ_n	Under each σ_n	Under all σ_n	Under each σ_n	Under all σ_n



1	Djg	100	16.699	15.003	16.699	14.090	0	-0.913
		200	15.563		15.642		0.079	
		300	15.110		14.216		-0.894	
		400	14.708		13.496		-1.212	
2	Ydg	100	29.683	27.954	27.474	26.778	-2.209	-1.176
		200	29.466		27.248		-2.218	
		300	27.923		26.870		-1.053	
		400	27.474		26.565		-0.909	
3	Dbz	100	38.660	32.822	37.235	32.293	-1.425	-0.529
		200	34.019		33.425		-0.594	
		300	33.024		32.619		-0.405	
		400	32.005		31.487		-0.518	

292

293 **5. Influence of the shearing rate on the residual friction angles according to soil properties**

294

295 Figure 9 depicts the relationships between residual friction angles as well as the difference in the residual friction
 296 angles and soil properties including liquid limit (LL), plasticity index (Ip) and clay fraction (CF) at two shearing rates.
 297 The residual friction angles at two shearing rates decreased nonlinearly with the increasing of the LL. As for the
 298 relationship between the ϕ_r and Ip, the ϕ_r under the low and high shearing rates decreases from about 32° to 15° with
 299 increasing the plasticity index from 11 to 16. With increasing of CF from 9% to 24%, the residual friction angles under
 300 low and high shearing rates were found to decrease. Interestingly, for Dbz and Ydg soils of which have similar
 301 percentage of clay fraction, the residual friction angles at both shearing rates varied. However, in the relationships
 302 between the difference in the residual friction angles and the soil properties, no clear correlations were found.

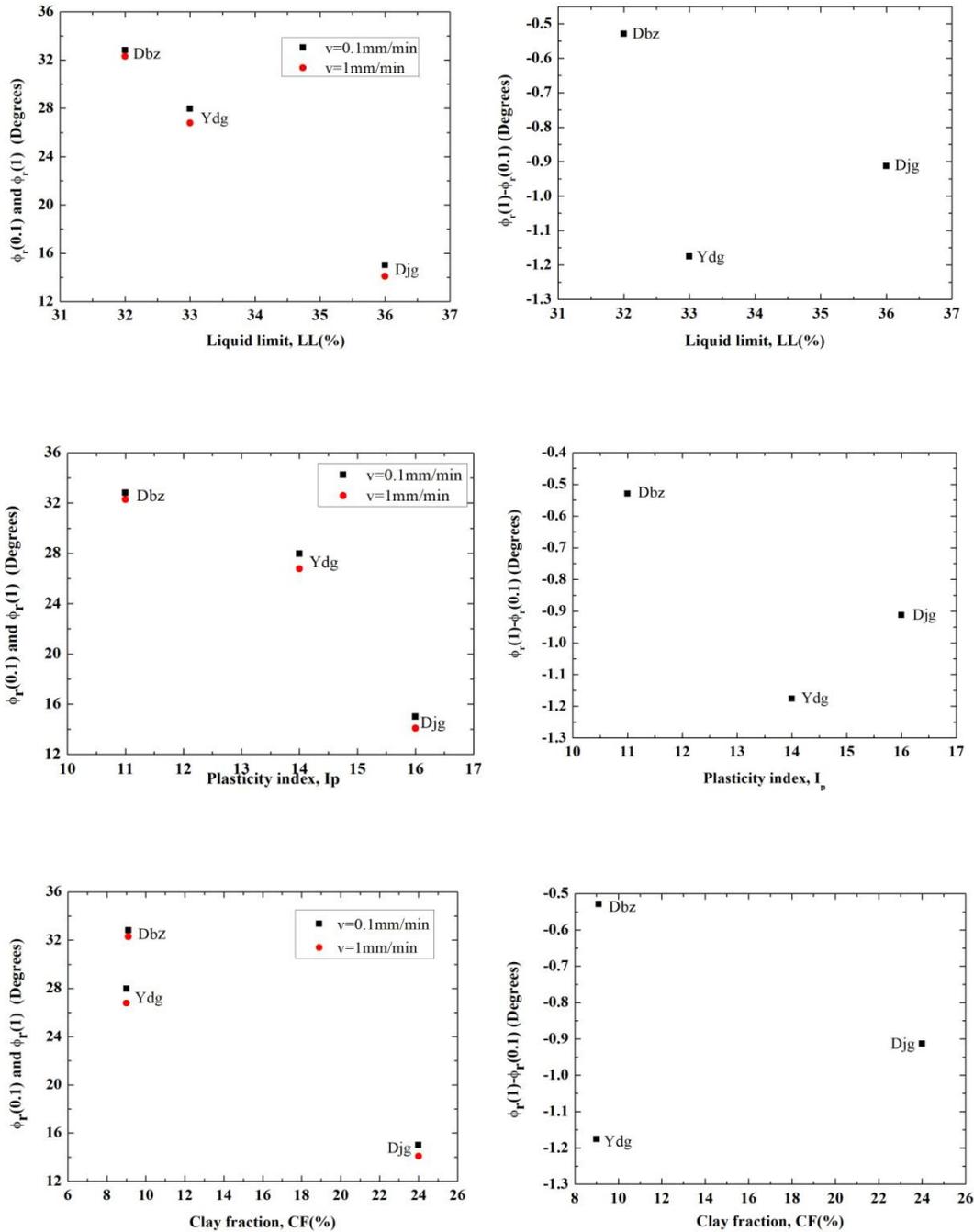


Figure 9. Relationships between residual shear parameter, the difference in residual shear parameter and the soil properties at two shearing rates



311 Conclusions

312 The shearing rate of slip zone soil of landslide may be changed after the occurrence of the first landslide activity, thus, the
313 residual shear strength of the slip zone soil could be changed accompanying this process. As for some precision engineering
314 which require high accuracy of the design parameters, the shearing rate effect on the residual shear strength should be fully
315 investigated. In this study, at the shearing rate of 0.1 mm/min and 1 mm/min, a series of ring shear tests under normal stress
316 ranged from 100kN/m² to 400kN/m² were performed on reconstituted slip zone soil samples obtained from three landslides
317 in loess area. The main results can be summarized in the following points:

- 318 (i) The shear displacement to achieve the residual stage for specimens with higher shearing rate is greater than that of
319 the lower rate.
- 320 (ii) As for slip zone soils in this study, specimens with lower shearing rate attained greater friction coefficients than
321 that with higher shearing rate.
- 322 (iii) At the two shearing rate (0.1 mm/min and 1 mm/min), the residual friction coefficient under the lower normal
323 stress was higher than that under the higher normal stress in all samples. In addition, there was a nonlinearly
324 decrease trend of the residual friction with the normal stress.
- 325 (iv) For slip zone soils in this study, the difference at the two shearing rate, $\phi_r(1) - \phi_r(0.1)$, under each normal stress
326 level were either negative or positive. However, under all normal stress, the difference at the two shearing rate
327 $\phi_r(1) - \phi_r(0.1)$ was found to be positive.
- 328 (v) The residual friction angles reduce with the increasing of shearing rate. Furthermore, the maximum magnitude of
329 the difference between the residual friction angle $\phi_r(1)$ and $\phi_r(0.1)$ was even obtained to be approximately 1.176°
330 in loess area.
- 331 (vi) The relationships between the ϕ_r under two shearing rates and soil properties including liquid limit and plasticity
332 index, demonstrated that the ϕ_r at both shearing rates decrease gradually with the increasing of LL and Ip.
333 However, no clear correlations between the difference in the ϕ_r at low and high shearing rates and the soil
334 properties were found.

335
336



337

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343



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